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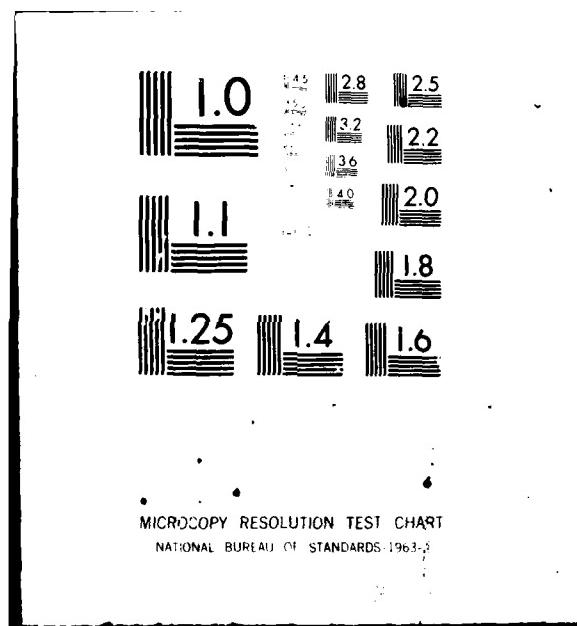
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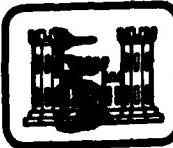
DOCUMENTATION FOR A TWO-LEVEL DYNAMIC THERMODYNAMIC SEA ICE MODEL

William D. Hibler III

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By



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20. ABSTRACT (Continue on reverse side if necessary and identify by block number) A discussion of the numerics and computer code for a two-level dynamic thermodynamic sea ice model is presented. For interested users a listing of the computer code and results from a 21-day test run are included as appendices. To a large degree this report is meant to serve as an extended appendix to an article by the author in the <i>Journal of Physical Oceanography</i> describing his model and a variety of simulation results. The model consists of a two-level ice thickness distribution coupled to the ice dynamics by a plastic rheology. In addition to the ice interaction the momentum balance includes nonlinear wind and water drag terms, Coriolis force, and inertial and momentum advection terms. The numerical scheme is formulated in an energy-conserving manner in a fixed Eulerian grid which allows simulation over unlimited time intervals. The momentum balance (including inertial terms) is numerically		

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20. Abstract (cont'd).

treated in a semi-implicit manner so that time steps of up to one day in length may be used if desired. The boundaries, grid size and time step magnitude are easily modified so that the model should have application to a variety of climate and forecasting problems.

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PREFACE

This report was prepared by William D. Hibler III, Research Physicist, of the Snow and Ice Branch, Research Division, U.S. Army Cold Regions Research and Engineering Laboratory. Funding for this research was provided by the Office of Naval Research, Code 461, Requisition No. N0001479MP90003, *Numerical Modeling of the Arctic Ice Cover and Ocean*, and by the National Aeronautics and Space Administration, Requisition No. C-27021-D, *Numerical Modeling of Sea Ice Dynamics and Ice Thickness Characteristics*.

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DOCUMENTATION FOR A TWO-LEVEL DYNAMIC THERMODYNAMIC SEA ICE MODEL

William D. Hibler III

INTRODUCTION

In a recent article Hibler (1979) formulated and tested a two-level numerical model for the simulation of sea ice circulation and thickness over a seasonal cycle. Although developed for long-term simulations, this model contains many features which may prove useful in sea ice forecasting. In addition the code may well have utility in studies of ice covered lakes.

The purpose of this report is to document the numerical procedure and computer code used in this model. To aid in using this code, a complete listing and a printout of a double precision test run are attached as Appendices.

In this documentation frequent reference will be made to the article in the *Journal of Physical Oceanography* that describes the model (Hibler 1979 hereafter referred to as JPO79). In addition to describing a variety of simulation results, this article contains a complete description of the model equations and a discussion of the numerical scheme and finite difference code. Consequently, various equations in JPO79 will frequently be directly referred to by number in the text of this report.

BRIEF DESCRIPTION OF MODEL

Model structure

The overall structure (see Fig. 1) essentially consists of three main components. The first is *momentum balance* which includes air and water stress, Coriolis force, internal ice stress, inertial forces and ocean tilt (sections 2-b-1 and 2-b-2 of JPO79). Nonlinear boundary layers for both the ocean-ice and air-ice surface traction are used. A key component of this momentum balance is the force due to internal ice stress (eqs 5 and 6, JPO79). This force is based on a constitutive law which relates the ice stress to the strain rate and ice strength. For this model, a viscous-plastic constitutive law is used. Rigid plastic behavior is approximated in this law by allowing the ice to flow in a plastic manner for normal strain rates and to creep in a linear viscous manner for small strain rates (eq 4-11 and Fig. 1 and 2 in JPO79).

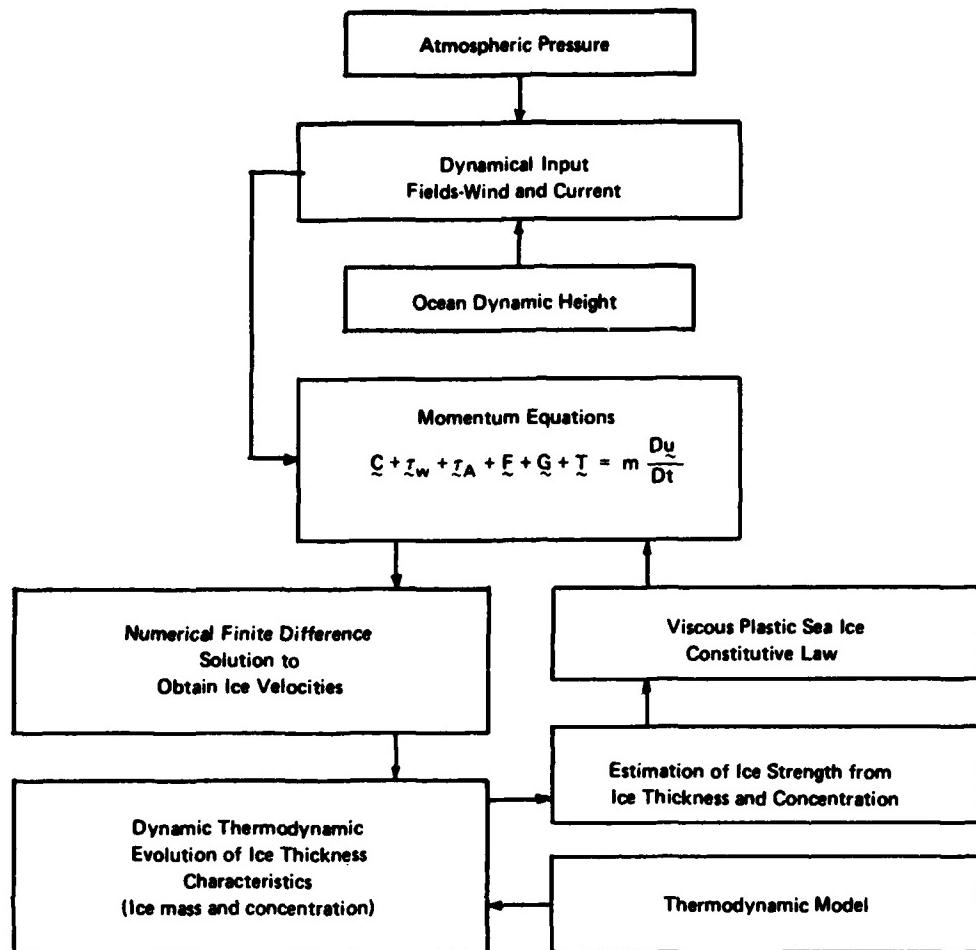
The treatment of ice as a viscous plastic fluid was largely motivated by the desire to avoid the complexities associated with elastic plastic rheologies (e.g. Pritchard 1975) while still retaining plastic behavior under flow.

The second major component of the model consists of *continuity equations* (section 2-b-3, JPO79) describing the evolution of the ice thickness characteristics. These equations are essentially a simplification to two levels of the Thorndike et al. (1975) ice thickness distribution model. In this particular code two categories of ice are assumed : thin ice (<0.5 m in thickness) and thick ice (>0.5 m). To keep track of these two categories, two parameters are calculated: the mean ice thickness per unit area h and the compactness A which is defined as the fraction of area covered by thick ice. The continuity equations describing the evolution of these parameters (eq 13-16, JPO79) also include thermodynamic terms. For testing the model geographically invariant growth rates as a function of thickness and time of year were used.

A final component is an *ice strength*. For this two-level case the strength is taken to depend linearly upon thickness, and exponentially upon compactness (eq 17, JPO79).

Numerical scheme

The coupled nonlinear equations of the model are treated as an initial value problem using energy conserving finite difference techniques. A fixed grid is used so that the set of equations may be integrated as long



- C Coriolis Force
 τ_w Water stress due to ice motion
 τ_A Air stress
 F Internal ice stress variation
 G Ocean currents
 T Ocean tilt
 $\frac{Du}{Dt}$ Ice acceleration and momentum advection
 m Ice mass per unit area

Figure 1. Flow chart for two-level dynamic thermodynamic sea ice model.

as desired. Under this initial value approach, values of all parameters at all grid points are required to start the integration, and boundary values of velocity are required thereafter. The basic forcing fields for the model consist of geostrophic wind fields and vertically integrated ocean currents. Both inertial and momentum advection terms are retained in the dynamical equations.

Because of the very strong ice interaction term, explicit integration of the momentum equations would force time steps to be extremely small (of the order of a few minutes). The ice thickness equations, on the other hand, can be explicitly integrated over time steps of several days. To avoid this severe time step limitation due to the ice interaction, the momentum equations

are integrated implicitly. This implicit integration does, however, require a relaxation solution of a set of simultaneous equations at each time step. (Details on the relaxation scheme are given in App. A, JPO79.)

Using this implicit scheme the inertial terms will normally be significant for short time steps (of the order of 1 hour or less) but insignificant for long time steps. Although not a formal stability requirement, it is wise to choose time steps that are small compared to the variability of the ice forcing because the nonlinear terms are treated in a semi-implicit manner in the momentum balance. The effect of different time step magnitudes on the momentum balance is discussed in Appendix B of JPO79. This appendix also contains some examples which may help the user decide on time step magnitudes for particular applications. Because of this implicit treatment of the momentum equation, the only formal stability requirement is a Courant-Friedrichs-Lowy condition for the advection terms in the thickness equations:

$$\Delta t < \Delta x [2(u^2 + v^2)]^{-\frac{1}{2}}.$$

A key feature of the numerical scheme is a staggered grid (Fig. 5, JPO79). This allows both the ice strength (i.e. the nonlinear viscosities) and the ice velocities to vary in space. Also this configuration greatly aids in constructing energy-conserving finite differences. The spatial finite differences used in the numerical scheme are described in some detail in Appendix A of JPO79. To a large degree this staggered grid is patterned after that employed in primitive equation numerical ocean models (e.g. Bryan 1969).

In the time-marching scheme (see eq 21-24, JPO79) ice thickness and velocity are temporally staggered. In particular, the thickness characteristics are considered to be defined at a temporal location halfway between the ice velocities. This scheme (formally called the forward-backward procedure) efficiently integrates coupled equations of this kind (e.g. Mesinger and Arakawa 1976).

As mentioned above, initial conditions at all points and ice velocities at the boundaries are thereafter required to initiate the integration of the system of equations forward in time. The most natural boundary condition is to take the ice velocity to be zero on the boundaries. This can be done either at a land boundary or at an ocean location where there is no ice. Note that the boundary condition does not affect the ice motion in such circumstances since in the absence of ice the strength is zero. More generally, as long as the ocean boundaries are removed from the ice edge, the coupled nature of the model will cause a natural ice edge boundary condition to be created (i.e. the ice strength drops to zero naturally near the edge). However, it is also

possible to form an "open" boundary condition by setting the strength equal to zero near a boundary. In JPO79 this type of open boundary condition is used at the Spitsbergen-Greenland passage to allow natural inflow or outflow. In the computer code three boundary meshes are used to define arbitrary boundaries and "open" boundaries. Islands are also allowed. Consequently by simply modifying these meshes highly irregular boundaries may be taken into account.

Because of the strong ice interaction (which in this model is dissipative in nature) the momentum equations are essentially parabolic in form and hence have few numerical instability problems over long-term integrations. (While there are few such numerical problems it should be emphasized that the dissipative ice interaction terms are highly nonlinear and can lead to unstable flow fields in the absence of water drag. Such a feature is a physical characteristic of plastic flow and not a numerical artifact). However, in principle it is possible for the ice interaction to be very small even though there may be a finite ice mass. Under this situation, the momentum advection terms could cause nonlinear instabilities. To ensure against such situations, the bulk viscosity parameter (and hence indirectly shear viscosity as well) is never allowed to drop below $4.0 \times 10^8 \text{ kg s}^{-1}$, a value which negligibly modifies the ice drift.

Nonlinear instabilities over long-term integrations can also arise from the nonlinear advection terms in the thickness continuity equations. To avoid this problem, small biharmonic and harmonic diffusion terms have been added to the continuity equation. These terms are grid-size dependent in the code and are automatically made smaller for finer mesh grids. These diffusion terms are treated in a mass-conserving manner at all boundaries (see App. A, JPO79). In practice these diffusive terms make only small contributions (in the JPO79 seasonal simulation the average diffusive flux was less than 3% of the average advective flux). It also appears that in the fully coupled model the diffusion terms might be dispensed with. Some sensitivity tests of the effects of diffusion are discussed in Appendix C of JPO79.

To integrate this model over very long time intervals requires a thermodynamic code specifying ice growth rates as a function of thickness and time of year. In the present code, spatially invariant growth rates (applicable to the Arctic Basin) from Thorndike et al. (1975) are used. For interested readers, time-independent computational procedures particularly useful for such growth calculations have been developed by Semtner (1976) for perennial multi-year ice, and Maykut (1978) for thin first-year ice.

Table 1. Schematic listing of model subroutines.

Level 1	Level 2	Level 3
	ADVECT (100-930) — DIFFUS (1190-1490)	
	BNDRY (940-1180)	
	RELAX (8470-10720)	FELLD (1500-1730) FELLD1 (1740-1990) FELLIP (2000-2330) XMAXM (12120-12230)
MAIN (4880-7810)	FORM (2500-3540) — PLAST (8000-8460) GROWTH (4100-4870) — GROATE (3810-4090) GEO (3550-3800) MEAN (7820-7990) FGROWP (2340-2490) XSUM (12240-12350)	

COMPUTER CODE

A complete test listing of the computer code is given in Appendix A. In this test code an idealized 5×6 grid is used. All arrays in this program are indexed beginning at the lower left-hand corner point (1, 1). Columns are denoted by constant x values, and rows by constant y values [i.e. the bottom row of an array consists of the points (1, 1), (2, 1) etc.]. In Appendix B double precision test run results of the code are presented.

In discussions of the code below, line numbers of specific statements will be given in parentheses. It is also noted that, while in principle the code is applicable to rectangular grid cells, it has only been tested for square grid cells ($D_x = D_y$).

Overall structure

The code is essentially divided into three levels. The first level consists of the main driving program (4880-7810) which sets up the boundaries, calls subroutines to numerically integrate the system of equations forward in time, and analyzes and stores results. (A description of the flow in the main program is given below.) The second level consists of various subroutines to numerically advance equations in time and to define various nonlinear parameters. These subroutines, in turn, call specialized subroutines at level 3 to, for example, estimate growth rates or determine nonlinear viscosities for plastic flow on a prescribed yield surface. Table 1 gives the schematic location; Table 2 briefly summarizes the function of each subroutine and Table 3 gives a brief description of the major variables. In all subroutines and the main driving program, a parameter statement is used to define dimension variables. Thus, by a simple editing command, array sizes can be modified everywhere in the code.

Note from the program listing in Appendix A that

the thickness and compactness grids are one unit larger in dimension in each direction than the velocity grid. With the staggered grid configuration this allows thickness values to be defined outside the velocity boundaries, a procedure which is convenient for second order differencing in the thickness equations. As mentioned earlier key features of the code are thickness and velocity masks which define the boundaries. These are defined by subroutine BNDRY. Specifically, UVM is a mask for the ice velocity (1 at calculated velocity locations, 0 at boundaries; islands are allowed), HEFFM is the thickness mask (1 inside boundaries, 0 outside), and OUT is an outflow or open boundary mask (same as HEFFM except that it is zero at grid cells where an open boundary is required). For the test case the masks (defined in a file called BNDDATA—see line 1005, Appendix A) create the grid configuration shown in Figure 2. At outflow grid cells the ice strength is set equal to zero and the ice mass removed. (All the mass removed is recorded in the main program.) Also, to approximately preserve second-order accuracy for advection into or out of an open cell, the mass and compactness in outflow grid cells are estimated by taking an average from all the grid cells adjacent to the open boundary (subroutine MEAN).

The input fields to the model consist of Geostrophic wind, Geostrophic Ocean currents (from which tilt is also calculated by the code) and growth rates. Other parameters are simulated (except for initial conditions and boundary values). Some description of the preparation of input fields from atmospheric pressure data and ocean currents is given in JPO79, section 3.

Appendix B gives the results of a 21-day test run of the complete code. For documentation purposes these test results are in double precision. In practice, however, double precision is not needed [except possibly, in the relaxation subroutine, subroutine RELAX, for very small grid sizes (say <25 km)].

Table 2. Brief description of model subroutines.

Name and Location	Function
MAIN (4880-7810)	Main driving program for model.
ADVECT (100-930)	Performs advection by explicit time stepping (modified Euler).
DIFFUS (1190-1490)	Performs diffusion by explicit forward time steps.
BNDRY (940-1180)	Sets up boundary masks (UVM, HEFFM, OUT).
RELAX (8470-10720)	Solves linearized momentum balance with spatially varying bulk and shear viscosities. Uses over-relaxation techniques.
FELLD (1500-1730)	
FELLD1 (1740-1990)	{ Calculate various finite differences for use in subroutine RELAX.
FELLIP (2000-2330)	
XMAXM (12120-12230)	Finds location of maximum value of a vector.
FORM (2500-3540)	Sets up forces, nonlinear water drag coefficients and nonlinear viscosities (for plastic flow) for use in each time step.
PLAST (8000-8460)	Calculates nonlinear viscosities based on plastic flow (specified by an elliptical yield curve) for subroutine FORM.
GROWTH (4100-4870)	Calculates changes in compactness A and mean ice thickness h due to growth, and redistribution by ridging.
GROATE (3810-4090)	Calculates growth rates for thin and thick ice in each grid cell.
GEO (3550-3800)	Reads in geostrophic wind data at each time step and ocean currents at the beginning of the model run.
MEAN (7820-7990)	Estimates thickness and compactness at "open boundary" grid cells based on adjacent values.
FGROWP (2340-2490)	Defines growth rates vs thickness. (In seasonal runs, this subroutine is not used. Instead a year-long matrix of growth rates is read in at line 5680 and daily values taken from this matrix at lines 6200-6230.)
XSUM (12240-12350)	Sums up values of a vector.

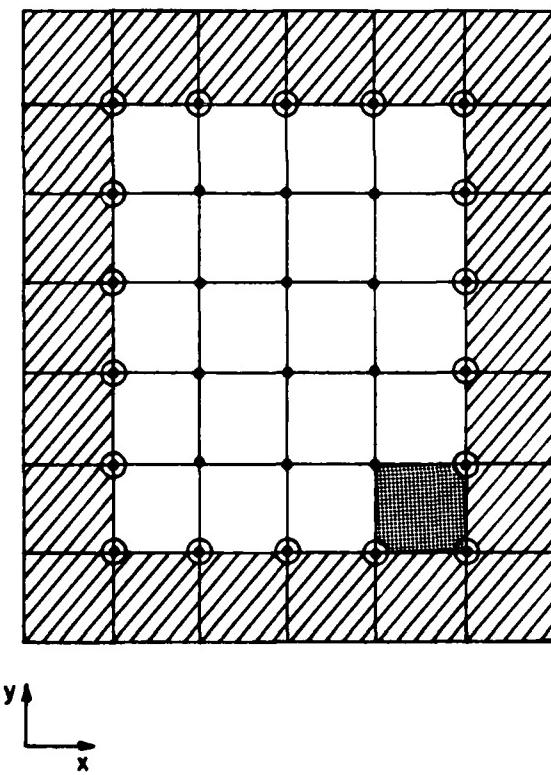


Figure 2. Boundary configuration for test grid. Circled grid points have zero velocities; shaded grid cells have zero thickness and compactness. Speckled grid cell is an "outflow" cell.

Description of flow in main driving program

To aid in understanding and using this code it is helpful to trace through the initialization and time-stepping procedure in the main driving program (lines 4880-7810). Such a procedure also gives some insight into various subtleties of the numerics as well as of the computer code. In this flow description, statements will be referred to by line number.

The first part of the main driving program (4880-6170) is concerned with initialization. Early in the code (5100-5220) constants are defined (1-day time steps, 125-km square grid, etc.) and various counters initialized (5230-5300). At lines 5310-5520 the boundary masks are read in (using subroutine BNDRY) and various sums over the boundaries performed to determine the number of grid cells covered by ice, etc. The program also prints out all the boundary masks to ensure that they are, in fact, the desired boundaries. At lines 5530-5640 the geostrophic currents are read in along with the first day's geostrophic wind data (using subroutine GEO). Note that a parameter WK (see line 5540) is used to cause GEO to either read in new currents or to use old ones. At line 5650-5690, a set of growth rates is read in (using subroutine FGROWP). In the main seasonal simulation in JPO79, FGROWP is ignored and a seasonal matrix is read off the tape at line 5680. However, for ice forecasting where the growth rates do not change very much, FGROWP with a data statement could have been used. FGROWP currently has January growth rates from Thorndike et al. (1975). Growth rates are entered

Table 3. Description of major variables (symbol in parentheses refers to variable symbol used in JPO79; e.g. h refers to ice thickness, A to compactness, etc.).

Symbol	Description	Location in staggered grid (see Fig. 5, JPO79)
<i>Arrays</i>		
UICE (u)	x component of ice velocity	corner
VICE (v)	y component of ice velocity	corner
HEFF (h)	mean ice thickness per unit area	center
AREA (A)	compactness (fraction of area covered by thick ice)	center
ETA (η)	nonlinear shear viscosity	center
ZETA (ζ)	nonlinear bulk viscosity	center
AMASS (m)	ice mass per unit area	corner
DRAGS (D_s)	symmetric nonlinear water drag component parameter plus inertial term	corner
DRAGA ($D_a + \lambda$)	antisymmetric nonlinear water drag plus Coriolis parameter	corner
GAIRX [U_g (x component)]	x component of geostrophic wind	corner
GAIRY [U_g (y component)]	y component of geostrophic wind	corner
GWATX [U_w (x component)]	x component of geostrophic ocean current	corner
GWATY [U_w (y component)]	y component of geostrophic ocean current	corner
FORCEX [$r_x + (\partial P / \partial x)$]	x component of external force plus ice pressure gradient	corner
FORCEY [$r_y + (\partial P / \partial y)$]	y component of external forces plus ice pressure gradient	corner
UVM	boundary mask for velocity field	corner
HEFFM	boundary mask for ice thickness field	center
OUT	boundary mask for thickness excluding "open boundaries"	center
HCORR	amount of ice needed to remove "negative ice thickness"	center
UICEC {	intermediate ice velocities for use in semi-implicit time step of momentum balance	corner
VICEC }	difference between HEFFM and OUT	center
OUT1	dummy variable for use in HMEAN	center
OUT2		center
UERR {	used to obtain velocity differences between time steps	corner
VERR }		corner
HDIFF	variable to keep track of diffusion	center
FHEFF [$f(h/A)$]	growth rate of thick ice	center
FO [$f(o)$]	growth rate of thin ice	center
STRESS (α)	xx , yy , and xy components of ice stress (mks units)	center
HDF	variable to aid in keeping track of diffusion	center
COR (λ)	Coriolis parameter	corner
DAIRN	nonlinear air stress coefficient	corner
($= \rho_a C_a U_g $)		corner
DWATN	nonlinear water stress coefficient	corner
($= \rho_w C_w U_w - u $)		corner
PRESS (P)	ice pressure or strength	center
ZMAX (ζ_{\max})	maximum allowable bulk viscosity value	center
ZMIN (ζ_{\min})	minimum allowable bulk viscosity value	center
GHEFF	average growth tendency	center
GAREA	area change tendency	center
E11 (ϵ_{xx})	xx strain rate component	center
E22 (ϵ_{yy})	yy strain rate component	center
E12 (ϵ_{xy})	xy strain rate component	center
<i>Vectors</i>		
FGROW (f)	growth rate of ice at 0.5-m interval [There are also a variety of vectors in the program which are equivalenced to arrays. This was done to facilitate sums and maximum point location which can be vectorized when vectors are used on the Texas Instruments Advanced Scientific Computer (ASC) for which this code was initially written.]	center

Table 3 (cont'd). Description of major variables (symbol in parentheses refers to variable symbol used in JPO79, e.g. h refers to ice thickness, A to compactness, etc.).

Symbol	Description
<i>Constants</i>	
DELTAT AND DELTAT	time step length in seconds
DELTAX	grid size in meters
DELTAY	grid size in meters
ERROR	accuracy of relaxation solution in m/s.
DIFF1 (D_1)	harmonic diffusion constant
A22	minimum allowable value of compactness
HO (h_0)	demarcation thickness between thin and thick ice
RHOICE (ρ_i)	ice density
FCOR (λ)	average Coriolis parameter
RHOAIR (ρ_a)	air density
SINWIN	sine of air turning angle
COSWIN	cosine of air turning angle
SINWAT	sine of ocean turning angle
COSWAT	cosine of ocean turning angle
GRAV (g)	acceleration of gravity
ECCEN (ϵ)	ratio of principal axes of plastic yield ellipse
PSTAR (P^*)	strength constant
THETA	weighting for momentum time stepping. THETA = 1.0 yields a backwards time step; THETA = 0.5 a Crank-Nicholson centered time step and 0.0 a forward explicit time step.

in units of centimeters per day and converted to meters of ice per second. Lines 5700-5930 initialize the ice thickness and velocity. [Note that the velocity, thickness and compactness functions are defined at three time levels, e.g. $H(l, j, K)$ $K = 1, 2, 3$. Of these values the $K = 1$ level is always the present value and the $K = 2$ level the previous time step.] As part of this initialization, line 5810 calls a subroutine to sum up the thicknesses at level 1 over the grid, excluding the "outflow" grid cells. This is accomplished by using the vector HEFF1 which is equivalence to the thickness array HEFF. (Note that the length of HEFF1 is such that only the first level of HEFF is included in HEFF1. This procedure has been followed on almost all the equivalence vectors.) In this initialization, the ice is set to 100% compactness with a uniform thickness of 3.2967m (= 3.0/density of ice).

The remainder of the initialization (5940-6160) is concerned with estimating the initial ice velocity field. (For long-term integrations this step is not critical.) To do this a spatially constant velocity is chosen, the ice mass is set equal to zero (effectively removing the inertial term from the relaxation solution), and the parameter THETA is set equal to 1. This procedure gives a backwards time step—see eqs 21 and 22, JPO79—so that the relaxation solution solves for the entire ice velocity rather than portions of it, as would happen in a centered time step, e.g. THETA = 0.5.

Having initialized the variables, the standard time-stepping procedure is begun at line 6170. After the conclusion of a given time step, the program returns to line 6180 (more precisely to FORTRAN statement number 100) and cycles through again. To initiate the integration, growth rates applicable to the day under consideration are defined from the growth rate matrix (line 6220). Since the present wind field is still valid, a new wind field is not called in until the end of the time step—to be mentioned later. Since only fixed January growth rates are used in the test case, line 6220 is commented. (Note that in this particular FORTRAN program comments are denoted by asterisks.) In lines 6240-6351 thicknesses and compactnesses at outflow points are estimated using subroutine MEAN. Basically this subroutine uses grid cells containing ice adjacent to the "open" boundaries to estimate ice thicknesses for use in advection. All the ice flowing into the "open" boundary cell is, however, explicitly accounted for. Part of this accounting is done at line 6351 where the amount of ice in the open cells (THEFF1) at the beginning of the time step is calculated.

At lines 6360-6550 the "predictor" portion of the momentum time step is performed (see eq 21, JPO79). A feature of this process is that at lines 6390-6420 ice velocities at the third level [e.g. UICE (l, j, 3)] and "centered" ice velocities [UICEC (l, j)] are set equal to the present ice velocities (level one) before calling

RELAX. This is done for two reasons: 1) the initial velocity "guess" to start the relaxation procedure is the ice velocity at the third level and 2) the momentum advection term is linearized by using UICEC (see eq 21 and App. A, JPO79). To carry out the predictor procedure, the time step is halved (line 6510) and FORM is called. FORM uses the present ice velocity to linearize the momentum equations as discussed earlier.

In lines 6560-6760 the main forward time step for the momentum balance is carried out (see eq 22, JPO79). The procedure here is the same as in the predictor step except that a full time step is used and the "predicted" velocity is employed to estimate the various nonlinear terms. This predictor corrector procedure approximately centers the nonlinear terms while still maintaining a backward time step for the "diffusion" terms.

Following the momentum time step, the thickness, growth and decay terms are explicitly stepped forward in time in lines 6770-6820. These operations correspond to the finite difference equations (eq 23-24) in JPO79. The advection subroutine (ADVECT) handles the dynamical portions of the compactness and thickness continuity equations, while the growth subroutine treats the thermodynamic terms. (The growth subroutine would be the logical place to insert a more complete heat budget subroutine.) As noted at line 6820, it is important to call the growth subroutine after the advection subroutine. This is because GROWTH also handles the mechanical redistribution by insisting that A be less than or equal to 1. GROWTH also corrects for problems such as negative ice thicknesses which may arise because centered differencing of second-order accuracy is used in the advection code (such negative values can be avoided by upstream differencing, but such schemes are excessively diffusive). Numerical studies of advection (Mahlman and Sinclair 1977) show that such negative value removal substantially aids in tracer accuracy provided that it is applied immediately as done here. Note that in these thickness time steps the latest values of the ice velocity are used. As mentioned earlier, this effectively means that the coupled thickness-momentum equations are staggered in time (see eq 21-24, JPO79).

Once the time stepping is completed, much of the remainder of the code is devoted to checking various sums over the array to ensure conservation and to keep track of the various contributions to the thickness changes. In lines 6830-6933, the outflow ice and various average thicknesses and growth sums are calculated. From 6940-7230 complete information over the grid (at an interval specified by lines 6950-6955) is printed out. In lines 7240-7420 velocity change information is printed out (over the same interval). In lines 7440-7510 simulated data for each time step are written to tape for later analysis (these write statements are

given comments in the test code). From lines 7511-7516 sum information is printed out (every time step). In lines 7520-7640 new winds are read in, and the wind data file is rewound if a whole year is over. Finally a check is made on the main time-step counter to see if the integration is finished. If not, the program cycles back to FORTRAN line 100. If finished, the program terminates after writing out some restart parameters (line 7790) for continuing the integrations. Some of the other termination steps after line 7650 are at present superfluous. They were initially inserted for an examination of "spin down" characteristics of the model.

Some comments on the relaxation subroutine

Since the relaxation code is the heart of the dynamic equations, some brief comments on this subroutine are added here. As presently written, the relaxation subroutine is unnecessarily complex. There are several reasons for this, mostly based on computational efficiency considerations. For one, the relaxation code was separately converted to single precision for the seasonal runs discussed in JPO79 to speed up the code. For this purpose all variables except those beginning in O were defined in single precision. As a result, the input double precision variables are converted to single precision at the beginning of the code by a variety of DO loops. Since such loops are vectorizable they take negligible time on the ASC. In addition the subroutines FELLD, FELLD1 and FELLIP were converted to single precision. While some changes may well be useful, caution should be exercised in any tinkering here. This is because double-precision and single precision variables are sometimes necessarily mixed in the code due to certain ASC compiler peculiarities.

A second reason for the complexity of the relaxation code, also related to computational efficiency, is that a large portion of the relaxation sweep was vectorized by calculating separately (in the subroutine FELLD1) as much of the finite difference code as can be performed over the whole array without replacement. While increasing the complexity of the code, this "separation" of the finite differences more than doubles the speed of the relaxation sweep on the ASC.

A third point is that the relaxation code was made general with respect to backward or forward time steps. In particular by adjusting THETA (input variable OTHETA) backwards, centered (Crank-Nicholson) or forward time steps may be used. To effect this generality, stresses at the previous time step are constructed by calling subroutine FELLIP. In general the backwards time step ($\text{THETA} = 1.0$) is most stable but as mentioned in JPO79 it does overdamp inertial waves. In the Crank-Nicholson scheme ($\text{THETA} = 0.5$, see, for

example, Richtmyer and Morton 1967) this does not occur. Consequently such a scheme is probably preferable at smaller time steps, say < 1 hour. At longer time steps, however, splitting problems have been found to occur with the centered scheme. These problems arise because the inertial term is the only term pulling alternating velocity variations together and this term is usually small at large time steps. Some brief discussion of this is given in section 2-C of JPO79.

A final point is that after 100 sweeps of over-relaxation, the subroutine returns to straight relaxation. This is because, while normally faster, the over-relaxation procedure can sometimes diverge.

CONCLUDING REMARKS

As presently configured, this sea ice model has direct application to long-term simulations. However, it also has many features which may well prove useful in ice forecasting. In addition, aspects of the code may aid in simulating plastic flow in geophysical media other than sea ice. In both such applications, modification of this code will probably be helpful. In these cases, features of the code either relevant or necessary for long integrations may be less important. In addition, certain characteristics which play no major role in long-term simulations may prove important in forecasting.

A feature important for the long-term studies but less important for forecasting is the presence of non-linear (and to a lesser degree linear) instabilities. The main examples of these are alternating grid point instabilities occurring in the thickness and compactness fields due to the nonlinear advection terms. To a lesser extent similar problems may arise in the momentum equation due to the momentum advection terms. Physically these instabilities arise due to the nonlinear advection terms causing a cascade of energy to higher wave numbers. Due to the finiteness of the grid, this energy tends to pile up at the Nyquist wave number, producing high wave number instabilities. Because of the compressible nature of sea ice flow, conventional energy-conserving techniques cannot always prevent this. In this code, this energy cascade is largely removed by biharmonic diffusion. This diffusion preferentially damps out the high wave number energy while little affecting the lower wave number. (In some sense this procedure can be thought of as a "poor man's spectral model.") In short-term integrations, however, such instabilities never have time to build up. Consequently, in such applications there appears to be no problem in removing all diffusion terms in the thickness equations. This removal will also cause some slight linear instability (since the advection terms are not perfectly centered in

time). However, over short time intervals these instabilities are minor. For similar reasons these considerations also apply to the minimum ice viscosity used in the momentum equations.

A converse feature, which is more critical in the short term, is the problem of initialization. For long-term simulations, the model can be integrated long enough so that it eventually "forgets" the initial conditions. For short-term forecasts, however, the initial conditions largely dominate the forecast. How to minimize the impact of poor initial conditions is an area needing considerable research. One approach may be to use observed forcing data to run the model for a few days prior to the forecast. Thus the model may generate some of its own initial conditions. It may also be that update data can be continually assimilated together with this "initialization" run.

Another difficulty results from specifying boundary conditions for localized regions. Persistence may be useful here. In any case a variety of these problems need numerical examination.

Finally, features of perhaps more general interest are some of the plastic characteristics of sea ice flow. In ice models to date many of these characteristics have been suppressed due to the damping nature of the oceanic boundary layer. In the absence of this damping, however, the plastic flow is inherently unstable and becomes ill-defined without inclusion of the inertial terms. In particular, in the absence of water drag, the system will begin to accelerate once the plastic yield is exceeded. This behavior has been verified directly with this code. (Initial results suggest, however, that this particular phenomenon is not critically dependent on time-step magnitude.) Consideration of spatially varying ocean coupling effects with the attendant buildup of pressure terms may well further complicate these plastic flow fields. Also on shorter time scales, inertial oscillation in both the ice and ocean may be important. Investigation of such effects is needed. In these investigations, aspects of the code described here may prove useful.

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APPENDIX A: LISTING OF COMPUTER CODE OF TWO-LEVEL MODEL.

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100      SUBROUTINE ADVECT(UICE1, VICE1, HEFF, DIFF1, LAD)
110 * SPACER
120      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
130      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
140      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
150      DIMENSION HEFF(NX1, NY1, 3), UICE1(NX, NY, 3), VICE1(NX, NY, 3)
160      &, UICE(NX, NY), VICE(NX, NY)
170      COMMON/STEP/DELTAT, DELTAX, DELTAY
180      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
185      COMMON/DIFF/HDIFF(NX1, NY1)
190 * NOW DECIDE IF BACKWARD EULER OR LEAPFROG
200      LL=LAD
210      IF(LL EQ 1) GO TO 100
220 * BACKWARD EULER
230      DELTT=DELTAT
240      K3=2
250      K2=2
260      GO TO 101
270 * LEAPFROG
280 100  DELTT=DELTAT*2.0
290      K3=3
300      K2=2
310 101  CONTINUE
320 * NOW REARRANGE H'S
330      DO 401 J=1, NY
340      DO 401 I=1, NX
350      UICE(I, J)=UICE1(I, J, 1)
360      VICE(I, J)=VICE1(I, J, 1)
370 401  CONTINUE
380      DO 200 J=1, NY1
390      DO 200 I=1, NX1
400      HEFF(I, J, 3)=HEFF(I, J, 2)
410      HEFF(I, J, 2)=HEFF(I, J, 1)
415      HDIFF(I, J)=0.0
420 200  CONTINUE
430 202  CONTINUE
440 * NOW GO THROUGH STANDARD CONSERVATIVE ADVECTION
450      DELTX=DELTAT/(4.0*DELTAX)
460      DELTY=DELTAT/(4.0*DELTAY)
470      DO 210 J=1, NYM1
480      DO 210 I=1, NXM1
490      HEFF(I+1, J+1, 1)=HEFF(I+1, J+1, K3)-DELTX*((HEFF(I+1, J+1, 2)+HEFF
500      &(I+2, J+1, 2))*(UICE(I+1, J+1)+UICE(I+1, J))-(HEFF(I+1, J+1, 2)+HEFF
510      &(I, J+1, 2))*(UICE(I, J+1)+UICE(I, J)))-DELTY*((HEFF(I+1, J+1, 2)+HEFF
520      &(I+1, J+2, 2))*(VICE(I, J+1)+VICE(I+1, J+1))-(HEFF(I+1, J+1, 2)+HEFF
530      &(I+1, J, 2))*(VICE(I, J)+VICE(I+1, J)))
540 210  CONTINUE
550 * NOW DECIDE IF DONE
560      IF(LL EQ 2) GO TO 99
570      IF(LL EQ 3) GO TO 89
580      GO TO 102
590 89   CONTINUE
600 * NOW FIX UP H(I, J, 2)
610      DO 88 J=1, NY1
620      DO 88 I=1, NX1
630      HEFF(I, J, 2)=HEFF(I, J, 3)
640 88   CONTINUE
650      GO TO 102
660 99   CONTINUE
670 * NOW DO BACKWARD EULER CORRECTION
680      DO 220 J=1, NY1
690      DO 220 I=1, NX1
700      HEFF(I, J, 3)=HEFF(I, J, 2)
710      HEFF(I, J, 2)=0.5*(HEFF(I, J, 1)+HEFF(I, J, 2))
720 220  CONTINUE
730      LL=3

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740      K3=3
750      GO TO 202
760 102  CONTINUE
770 * NOW DO DIFFUSION ON H(I,J,K3)
780      DO 240 KD=1,2
790      GO TO (241,242),KD
800 241  CONTINUE
810      CALL DIFFUS(VICE, VICE, HEFF, DIFF1, DELTT)
820      GO TO 243
830 242  CONTINUE
840      DIFF2=-(DELTAX**2)/DELT
850      CALL DIFFUS(VICE, VICE, HEFF, DIFF2, DELTT)
860 243  CONTINUE
870      DO 330 J=1,NY1
880      DO 330 I=1,NX1
890      HEFF(I,J,1)=(HEFF(I,J,1)+HEFF(I,J,3))*HEFFM(I,J)
895      HDIFF(I,J)=HDIFF(I,J)+HEFF(I,J,3)*HEFFM(I,J)
900 330  CONTINUE
910 240  CONTINUE
920      RETURN
930      END
940      SUBROUTINE BNDRY
950 * SUBROUTINE SETS UP BOUNDARY MASK
960      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
970      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
980      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
990      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
1000      COMMON/OUTFL/OUT(NX1, NY1)
1005      OPEN(UNIT=1, FILE="BNDATA")
1010 * READ IN VELOCITY MASK
1020      READ(UNIT=1, FMT=9) ((UVM(I,J), I=1, NX), J=1, NY)

1030 9   FORMAT(30G1. 0)
1040      READ(UNIT=1, FMT=8) ((HEFFM(I,J), I=1, NX1), J=1, NY1)
1050      READ(UNIT=1, FMT=8) ((OUT(I,J), I=1, NX1), J=1, NY1)
1060 8   FORMAT(42G1. 0)
1170      RETURN
1180      END
1190      SUBROUTINE DIFFUS(VICE, VICE, HEFF, DIFF1, DELTT)
1200 * SPACER
1210      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
1220      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
1230      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
1240      DIMENSION HEFF(NX1, NY1, 3), UICE(NX, NY), VICE(NX, NY)
1250      &, HEFF1(NX1, NY1)
1260      COMMON/STEP/DELTAT, DELTAX, DELTAY
1270      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
1280 * SUBROUTINE DIFUSES HEFF, MULTIPLIES BY DELT, AND PUTS RESULTS IN HEFF
1290 * NOW ZERO OUT HEFF1
1300      DO 210 J=1, NY1
1310      DO 210 I=1, NX1
1320      HEFF1(I,J)=0. 0
1330 210  CONTINUE
1340 * NOW DO DIFFUSION
1350      DELTXX=DELT*DIFF1/(DELTAX**2)
1360      DELTYY=DELT*DIFF1/(DELTAY**2)
1370      DO 220 J=2, NY
1380      DO 220 I=2, NX
1390      HEFF1(I,J)=DELTXX*((HEFF(I+1,J,3)-HEFF(I,J,3))*HEFFM(I+1,J)
1400      & -(HEFF(I,J,3)-HEFF(I-1,J,3))*HEFFM(I-1,J))
1410      & +DELTYY*((HEFF(I,J+1,3)-HEFF(I,J,3))*HEFFM(I,J+1)
1420      & -(HEFF(I,J,3)-HEFF(I,J-1,3))*HEFFM(I,J-1))
1430 220  CONTINUE
1440      DO 260 J=1, NY1
1450      DO 260 I=1, NX1
1460      HEFF(I,J,3)=HEFF1(I,J)

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1470 260 CONTINUE
1480 RETURN
1490 END
2340 SUBROUTINE FGROWP
2350 * SUBROUTINE TO SET UP CANONICAL GROWTH VALUES
2360 IMPLICIT DOUBLE PRECISION (A-H,D-Z)
2370 PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
&, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
2380 DIMENSION FGGOW(21)
2390 COMMON/GROWTH/FGROW(21)
2400 DATA FGGOW/12 09D+00, 1 95D+00, . 46D+00, . 37D+00, . 31D+00, . 27D+00
2410 &, . 21D+00, . 14D+00, . 09D+00, . 03D+00, 0 0, - . 02D+00, - . 03D+00, - . 04D+00
2420 &, - . 05D+00, - . 05D+00, - . 06D+00, - . 06D+00, - . 06D+00, - . 06D+00, - . 06D+00/
2430 DO 101 I=1, 21
2440 FGROW(I)=FGGOW(I)/(8. 64D+06)
2450 PRINT 1, I, FGROW(I)
2460 101 CONTINUE
2470 1 FORMAT(1X, I5, G20 12)
2480 RETURN
2490 END
2500 SUBROUTINE FORM(VICE, VICE, ETA, ZETA, DRAGS
&, DRAGA, GAIRX, GAIRY, GWATX, GWATY, FORCEX, FORCEY, HEFF, AMASS, AREA)
2510 IMPLICIT DOUBLE PRECISION (A-H,D-Z)
2520 PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
&, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
2530 * PROGRAM FORMS BASIC INPUT PARAMETERS FOR RELAXATION
2540 DIMENSION VICE(NX, NY, 3), VICE(NX, NY, 3), ETA(NX1, NY1)
&, ZETA(NX1, NY1), DRAGS(NX, NY), DRAGA(NX, NY), GAIRX(NX, NY)
&, GAIRY(NX, NY), GWATX(NX, NY), GWATY(NX, NY), HEFF(NX1, NY1, 3),
&, COR(NX, NY), FORCEX(NX, NY), FORCEY(NX, NY), AMASS(NX, NY)
2550 &, DAIRN(NX, NY), DWATN(NX, NY), AREA(NX1, NY1, 3), PRESS(NX1, NY1)
&, ZMAX(NX1, NY1), ZMIN(NX1, NY1)
2560 COMMON/STEP/DELTAT, DELTAX, DELTAY
2570 COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
2580 COMMON/OUTFLD/OUT(NX1, NY1)
2590 * FIRST SET UP BASIC CONSTANTS
2600 DWAT=0. 59D+00
2610 DAIR=0. 01462D+00
2620 RHOICE=0. 91D+03
2630 FCOR=1. 46D-04
2640 RHOAIR=1. 3D+00
2650 SINWIN=0. 4226D+00
2660 COSWIN=0. 9063D+00
2670 SINWAT=0. 4226D+00
2680 COSWAT=0. 9063D+00
2690 GRAV=9. 832D+00
2700 ECCEN=2. 0
2710 * 20 DEG GIVES SIN EQUAL TO . 34202
2720 * NOW SET UP MASS PER UNIT AREA AND CORIOLIS TERM
2730 DO 101 J=1, NY
2740 DO 101 I=1, NX
2750 AMASS(I, J)=RHOICE*0. 25*(HEFF(I, J, 1)
& +HEFF(I+1, J, 1)+HEFF(I, J+1, 1)+HEFF(I+1, J+1, 1))
2760 COR(I, J)=AMASS(I, J)*FCOR
2770 101 CONTINUE
2780 * NOW SET UP NON LINEAR WIND AND WATER DRAG
2790 DO 99 J=1, NY
2800 DO 99 I=1, NX
2810 DAIRN(I, J)=RHOAIR*. 0012D+00*DSQRT(GAIRX(I, J)**2+GAIRY(I, J)**2)
2820 DWATN(I, J)=5. 5D+00*DSQRT((VICE(I, J, 1)-GWATX(I, J))**2
& +(VICE(I, J, 1)-GWATY(I, J))**2)
2830 99 CONTINUE
2840 * NOW SET UP SYMMETRIC DRAG
2850 DO 102 J=1, NY
2860 DO 102 I=1, NX
2870 DRAQS(I, J)=DWATN(I, J)*COSWAT

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2950 102 CONTINUE
2960 * NOW SET UP ANTI SYMMETRIC DRAG PLUS CORIOLIS
2970     DO 103 J=1,NY
2980     DO 103 I=1,NX
2990     DRAGA(I,J)=DWATN(I,J)*SINWAT+COR(I,J)
3000 103 CONTINUE
3010 * NOW SET UP FORCING FIELD
3020     DO 104 J=1,NY
3030     DO 104 I=1,NX
3040 * FIRST DO WIND
3050     FORCEX(I,J)=DAIRN(I,J)*(COSWIN*GAIRX(I,J)
3060     & -SINWIN*GAIRY(I,J))
3070     FORCEY(I,J)=DAIRN(I,J)*(SINWIN*GAIRX(I,J)
3080     & +COSWIN*GAIRY(I,J))
3090 * NOW ADD IN CURRENT FORCE
3100     FORCEX(I,J)=FORCEX(I,J)+DWATN(I,J)*(COSWAT*GWATX(I,J)
3110     & -SINWAT*GWATY(I,J))
3120     FORCEY(I,J)=FORCEY(I,J)+DWATN(I,J)*(SINWAT*GWATX(I,J)
3130     & +COSWAT*GWATY(I,J))
3140 * NOW ADD IN TILT
3150     FORCEX(I,J)=FORCEX(I,J)-COR(I,J)*GWATY(I,J)
3160     FORCEY(I,J)=FORCEY(I,J)+COR(I,J)*GWATX(I,J)
3170 104 CONTINUE
3180 * NOW SET UP ICE PRESSURE AND VISCOSITIES
3190 * FIRST SET UP CONSTANTS
3200     PSTAR=5.0D+03
3210     EPSIL=5.787D-08
3220 * 5.787D-08 IS 0.5 PER CENT PER DAY STRAIN RATE
3230     ZETAC=0.15
3240     ETAC=ZETAC/4.0
3250 * NOW SET UP VALUES
3260     DO 105 J=1,NY1
3270     DO 105 I=1,NX1
3280     PRESS(I,J)=PSTAR*HEFF(I,J,1)*DEXP(-20.0*(1.0-AREA(I,J,1)))
3290     ZMAX(I,J)=(5.0D+12/(2.0D+04))*PRESS(I,J)
3300     ZMIN(I,J)=4.0D+08
3310 *     ZETA(I,J)=ZETAC*PRESS(I,J)/EPSIL
3320 *     ETA(I,J)=ETAC*PRESS(I,J)/EPSIL
3330 *     ZETA(I,J)=DMAX1(4.0D+08,ZETA(I,J))
3340 *     ETA(I,J)=DMAX1(1.0D+08,ETA(I,J))
3350 105 CONTINUE
3360     CALL PLAST(VICE, VICE, PRESS, ETA, ZETA, ECCEN, ZMAX, ZMIN)
3370 * NOW SET VISCOSITIES AND PRESSURE EQUAL TO ZERO AT OUTFLOW PTS
3380     DO 106 J=1,NY1
3390     DO 106 I=1,NX1
3400     PRESS(I,J)=PRESS(I,J)*OUT(I,J)
3410     ETA(I,J)=ETA(I,J)*OUT(I,J)
3420     ZETA(I,J)=ZETA(I,J)*OUT(I,J)
3430 106 CONTINUE
3440 * NOW CALCULATE PRESSURE FORCE AND ADD TO EXTERNAL FORCE
3450     DO 107 J=1,NY
3460     DO 107 I=1,NX
3470     FORCEX(I,J)=FORCEX(I,J)-(0.25/DELTAX)*
3480     & (PRESS(I+1,J)+PRESS(I+1,J+1)-PRESS(I,J)-PRESS(I,J+1))
3490     FORCEY(I,J)=FORCEY(I,J)-(0.25/DELTAY)*
3500     & (PRESS(I,J+1)+PRESS(I+1,J+1)-PRESS(I,J)-PRESS(I+1,J))
3510 * NOW PUT IN MINIMAL MASS FOR TIME STEPPING CALCULATIONS
3520 107 CONTINUE
3530     RETURN
3540     END
3550     SUBROUTINE GEO(GAIRX, GAIRY, GWATX, GWATY, WK)
3560 * IF WK LT 0 THEN READ IN CURRENTS, IF WK GT 100 THEN ZERO OUT CURRENTS
3570     IMPLICIT DOUBLE PRECISION (A-H,O-Z)
3580     PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
3590     &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY

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3600      DIMENSION GAIWX(NX, NY), GAIY(NX, NY), GWATX(NX, NY), GWATY(NX, NY)
3610 * FIRST READ IN WIND
3620      DO 101 J=1, NY
3630      DO 101 I=1, NX
3640      GAIWX(I, J)=3.0D+00+0.1D+00*(I+J)
3650      GAIY(I, J)=-4.0D+00+0.2D+00*(I+J)
3660      GWATX(I, J)=0.01D+00-0.001D+00*(I-J)
3670      GWATY(I, J)=0.05D+00+0.001D+00*(I+J)
3680 101  CONTINUE
3790      RETURN
3800      END
3810      SUBROUTINE GROATE(H, HO, FH, FO, AREA)
3820 * SUBROUTINE TO OBTAIN THICK AND THIN GROWTH RATES
3830      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
3840      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
3850      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
3860      DIMENSION H(NX1, NY1), FO(NX1, NY1), FH(NX1, NY1)
3870      &, AREA(NX1, NY1, 3)
3880      COMMON/GROWTH/FGROW(21)
3890      ONE=1.0
3900      ZERO=0.0
3910      S2=HO+1.0
3920      M2=IDINT(S2)
3930      DO 101 J=1, NY1
3940      DO 101 I=1, NX1
3950      S1=H(I, J)/0.5+1.0
3960      S1=DMIN1(S1, 2.0D+01)
3970      M1=IDINT(S1)
3980      FO(I, J)=(FGROW(M2+1)-FGROW(M2))*(S2-DFLOAT(M2))+FGROW(M2)
3990 * THIN ICE IS ALWAYS TAKEN TO BE ZERO THICKNESS
4000 201  FO(I, J)=FGROW(1)
4010 202  CONTINUE
4020      FH(I, J)=(FGROW(M1+1)-FGROW(M1))*(S1-DFLOAT(M1))+FGROW(M1)
4030 101  CONTINUE
4040      DO 102 J=1, NY1
4050      DO 102 I=1, NX1
4060      FH(I, J)=FH(I, J)*AREA(I, J, 2)+FO(I, J)*(1.0-AREA(I, J, 2))
4070 102  CONTINUE
4080      RETURN
4090      END
4100      SUBROUTINE GROWTH(HEFF, AREA, FH, HO, A22)
4110 * SUBROUTINE TO CALCULATE CHANGE OF PARAMETERS DUE TO GROWTH
4120      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
4130      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
4140      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
4150 * THIS SUBROUTINE MUST BE CALLED AFTER ADVECTION
4160 * FH IS ADDITION MELTING THAT MUST BE DONE TO BALANCE MIXED LAYER
4170      DIMENSION HEFF(NX1, NY1, 3), AREA(NX1, NY1, 3), GHEFF(NX1, NY1)
4180      &, GAREA(NX1, NY1), H(NX1, NY1), FH(NX1, NY1)
4190      COMMON/RATE/FHEFF(NX1, NY1), FO(NX1, NY1)
4200      COMMON/STEP/DELTAT, DELTAX, DELTAY
4210      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
4220      COMMON/OUTFLO/OUT(NX1, NY1)
4230      ZERO=0.0
4240      ONE=1.0
4250 * FIRST SOLVE FOR THICKNESS OF THICK ICE
4260      DO 101 J=1, NY1
4270      DO 101 I=1, NX1
4280      AREA(I, J, 2)=DMAX1(A22, AREA(I, J, 2))
4290      H(I, J)=HEFF(I, J, 2)/AREA(I, J, 2)
4300 101  CONTINUE
4310 * GROWTH SUBROUTINE CALCULATES TOTAL GROWTH TENDENCIES INCLUDING SNOWFAL
4320      CALL GROATE(H, HO, FHEFF, FO, AREA)
4330 1      FORMAT(1X, 11G11.5)
4340 * NOW CALCULATE CORRECTED GROWTH
4350      DO 201 J=1, NY1

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4360      DO 201 I=1,NX1
4370      GHEFF(I,J)=-DELTAT*FHEFF(I,J)
4380      GAREA(I,J)=DELTAT*FO(I,J)
4390 201  CONTINUE
4400      DO 102 J=1,NY1
4410      DO 102 I=1,NX1
4420      GHEFF(I,J)=-1.0*DMIN1(HEFF(I,J,1),GHEFF(I,J))
4430      GAREA(I,J)=DMAX1(ZERO,GAREA(I,J))
4440 102  CONTINUE
4450      DO 103 J=1,NY1
4460      DO 103 I=1,NX1
4470      FH(I,J)=DMIN1(ZERO,GHEFF(I,J))
4480 103  CONTINUE
4490 * NOW DO AREA CHANGE
4500      DO 104 J=1,NY1
4510      DO 104 I=1,NX1
4520      GAREA(I,J)=2.0*(1.0-AREA(I,J,2))*GAREA(I,J)/HD
4530      &+.5*FH(I,J)*AREA(I,J,2)/(HEFF(I,J,1)+.00001)
4540 104  CONTINUE
4550 * NOW CORRECT AREA AND HEFF(I,J,1)
4560      DO 105 J=1,NY1
4570      DO 105 I=1,NX1
4580      AREA(I,J,1)=AREA(I,J,1)+GAREA(I,J)
4590      HEFF(I,J,1)=HEFF(I,J,1)+GHEFF(I,J)*OUT(I,J)
4600 105  CONTINUE
4610 * NOW ZERO OUTSIDE POINTS
4620      DO 109 J=1,NY1
4630      DO 109 I=1,NX1
4640      AREA(I,J,1)=AREA(I,J,1)*HEFFM(I,J)
4650      HEFF(I,J,1)=HEFF(I,J,1)*HEFFM(I,J)
4660 109  CONTINUE
4670 * NOW SET AREA(I,J,1)=0 WHERE NO ICE IS
4680      DO 110 J=1,NY1
4690      DO 110 I=1,NX1
4700      AREA(I,J,1)=DMIN1(AREA(I,J,1),HEFF(I,J,1)/.0001)
4710 110  CONTINUE
4720 * NOW TRUNCATE AREA
4730      DO 106 J=1,NY1
4740      DO 106 I=1,NX1
4750      AREA(I,J,1)=DMIN1(ONE,AREA(I,J,1))
4760 106  CONTINUE
4770      DO 107 J=1,NY1
4780      DO 107 I=1,NX1
4790      AREA(I,J,1)=DMAX1(A22,AREA(I,J,1))

4800 107  CONTINUE
4810 * NOW CALCULATE ADDITIONAL ICE TO BE MELTED FOR MIXED LAYER BALANCE
4820      DO 108 J=1,NY1
4830      DO 108 I=1,NX1
4840      FH(I,J)=GHEFF(I,J)-DELTAT*FHEFF(I,J)
4845 * NOW STORE COMMON GROWTH RATE
4846      FHEFF(I,J)=GHEFF(I,J)
4850 108  CONTINUE
4860      RETURN
4870      END
4880 * MAIN DRIVING PROGRAM FOR VISCOUS PLASTIC SEA ICE MODEL
4890 IMPLICIT DOUBLE PRECISION (A-H,O-Z)
4900 * SPACER
4910      PARAMETER NX=5,NY=6,N3=(NX+1)*(NY+1),NX1=NX+1
4920      &,NY1=NY+1,NXM1=NX-1,NYM1=NY-1,N4=NX*NY
4930      DIMENSION UIICE(NX,NY,3),VICE(NX,NY,3),ETA(NX1,NY1)
4940      &,ZETA(NX1,NY1),DRAGS(NX,NY),DRAGA(NX,NY),QAIRX(NX,NY)
4950      &,QAIRY(NX,NY),QWATX(NX,NY),QWATY(NX,NY),HEFF(NX1,NY1,3)
4960      &,AMASS(NX,NY),FORCEX(NX,NY),FORCEY(NX,NY),AREA(NX1,NY1,3)
4970      &,HCORR(NX1,NY1),HEFF1(N3)
4980      &,UICEC(NX,NY),VICEC(NX,NY)
4990      &,OUT1(NX1,NY1),OUT2(NX1,NY1),UERR(NX,NY),VERR(NX,NY),UERRV(N4)

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5000      &, VERRV(N4)
5010      &, FGRW(365, 21), HDIFF(NX1, NY1), FHE1(N3), GR(N3), AREA1(N3)
5020      COMMON/RATE/FHEFF(NX1, NY1), FO(NX1, NY1)
5030      COMMON/STEP/DELTAT, DELTAX, DELTAY
5040      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
5050      COMMON/OUTFLO/OUT(NX1, NY1)
5060      COMMON/GROWTH/FGROW(21)
5070      COMMON/STRESS/STRESS(NX1, NY1, 3)
5072      COMMON/DIFF/HDF(NX1, NY1)
5073      EQUIVALENCE(HDIFF, GR)
5074      EQUIVALENCE(FHEFF, FHE1)
5075      EQUIVALENCE(FHEFF, FHE1)
5076      EQUIVALENCE(AREA, AREA1)
5080      EQUIVALENCE(HEFF, HEFF1)
5090      EQUIVALENCE(UERR, UERRV), (VERR, VERRV)
5100 * NOW DECIDE ON BASIC PARAMETERS
5110      DELTT=86400. 0
5120      DELTAT=86400. 0
5130      DELTAX=1. 25D+05
5140      DELTAY=1. 25D+05
5150      ERROR= 000001D+00
5160      ERROR=5. 0D+00*ERROR
5162      FHSUM1=0. 0
5163      GRSUM1=0. 0
5164      ARSUM1=0. 0
5170      DIFF1=. 002D+00*DELTAX
5180      DIFF1=2. 0*DIFF1
5190      A22=0. 15D+00
5200      HO=0. 5
5210 * DOUBLE HO BECAUSE OF MOD IN GROWTH
5220      HO=2. 0*HO
5230      LAD=2
5240      KSTOP=1

5250 * NOW INITIALIZE COUNTER
5260      ICOUNT=0
5270      TOUT=0. 0
5280      HESUM=0. 0
5290      UVSUM=0. 0
5300      HHESUM=0. 0
5310 * NOW DEFINE BOUNDARIES
5320      CALL BNDRY
5330 *      WRITE(20, ((OUT(I, J), I=1, NX1), J=1, NY1)
5340      DO 122 J=1, NY1
5350      DO 122 I=1, NX1
5360      OUT1(I, J)=HEFFM(I, J)-OUT(I, J)
5370      HESUM=HESUM+OUT(I, J)
5380      HHESUM=HHESUM+HEFFM(I, J)
5390      122 CONTINUE
5400      DO 125 J=1, NY
5410      DO 125 I=1, NX
5420      UVSUM=UVSUM+UVM(I, J)
5430      125 CONTINUE
5440      PRINT 3, ((HEFFM(I, J), I=1, NX), J=1, NY)
5450      PRINT 4
5460      PRINT 3, ((UVM(I, J), I=1, NX), J=1, NY)
5470      PRINT 4
5480      PRINT 3, ((OUT(I, J), I=1, NX), J=1, NY)
5490      PRINT 4
5500      PRINT 6, HESUM, UVSUM
5510      PRINT 6, HHESUM, UVSUM
5520      6 FORMAT(1X, 'HESUM IS', Q15. 6, 'UVSUM IS', Q15. 7)
5530 * NOW READ IN GEOSTROPHIC WIND AND CURRENTS
5540 * IF WK LT 0 THEN NEW CURRENTS READ IN, IF WK GT 100 THEN CUR ZERODED
5550 *      31 IS WIND FILE
5560 *      30 IS CURRENT FILE
5570      WK=-1. 0D+00

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5580      CALL CED(GAIRX, GAIRY, GWATX, GWATY, WK)
5590      PRINT 3, ((GAIRX(I, J), I=1, NX), J=1, NY)
5600      PRINT 4
5610      PRINT 3, ((GAIRY(I, J), I=1, NX), J=1, NY)
5620      PRINT 4
5630      PRINT 3, ((GWATX(I, J), I=1, NX), J=1, NY)
5640      PRINT 4
5650 * NOW SET UP GROWTH RATES
5660      CALL FGROWP
5670 * FILE 32 CONTAINS GROWTH RATES
5680 * READ(32) FGROW
5690 * NOW INITIALIZE SYSTEM
5700 * FIRST GUESS AT INITIAL HEFF AND AREA
5710      DO 101 J=1, NY1
5720      DO 101 I=1, NX1
5725      HDIFF(I, J)=0.0
5730      HEFF(I, J, 3)=0.0
5740      HEFF(I, J, 2)=0.0
5750      AREA(I, J, 2)=1.0
5760      AREA(I, J, 3)=1.0
5770      HEFF(I, J, 1)=(3.0D+00)/0.91D+00
5780      HEFF(I, J, 1)=HEFF(I, J, 1)*OUT(I, J)
5790      AREA(I, J, 1)=1.0
5800 101  CONTINUE
5810      CALL XSUM(HEFF1, THEFF)
5820 * NOW ZERO OUT VELOCITY
5830      DO 102 J=1, NY
5840      DO 102 I=1, NX
5850      UICE(I, J, 3)=0.0
5860      UICE(I, J, 2)=0.0
5870      UICE(I, J, 1)=0.0
5880      VICE(I, J, 3)=0.0
5890      VICE(I, J, 2)=0.0
5900      VICE(I, J, 1)=0.0
5910      VICEC(I, J)=0.0
5920      VICEC(I, J)=0.0
5930 102  CONTINUE
5940 * NOW GET FIRST VALUE OF U AND V
5950      THETA=1.0
5960      CALL FORM(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, GAIRX, GAIRY, GWATX, GWATY
5970      &, FORCEX, FORCEY, HEFF, AMASS, AREA)
5980      PRINT 3, ((FORCEX(I, J), I=1, NX), J=1, NY)
5990      PRINT 4
6000      PRINT 3, ((FORCEY(I, J), I=1, NX), J=1, NY)
6010 * NOW SET AMASS=0 AND DEFINE INITIAL VISCOSITY
6020      DO 103 J=1, NY
6030      DO 103 I=1, NX
6040      AMASS(I, J)=0.0
6050      ZETA(I, J)=HEFF(I, J, 1)*(1.0D+11)
6060      ETA(I, J)=ZETA(I, J)/4.0
6070 103  CONTINUE
6080      CALL RELAX(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, AMASS, FORCEX, FORCEY
6090      &, ERROR, THETA, UICEC, VICEC)
6100      PRINT 4
6110      PRINT 4
6120      PRINT 3, ((UICE(I, J, 1), I=1, NX), J=1, NY)
6130      PRINT 4
6140      PRINT 3, ((VICE(I, J, 1), I=1, NX), J=1, NY)
6150      PRINT 4
6160      PRINT 4
6170 * NOW START STANDARD PREDICTOR CORRECTOR PROCEDURE
6180 100  CONTINUE
6190 * NOW PUT GROWTH RATES IN
6200      KQRO=MOD(ICOUNT, 365)+1
6210 *      DO 141 KQQ=1, 21
6220 *      FGROW(KQQ)=FGROW(KQRO, KQQ)/(8.64D+06)

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6230 *141 CONTINUE
6240 * NOW SET UP REASONABLE OUTFLOW THICKNESSES
6250 CALL MEAN(HEFF, OUT2)
6260 DO 123 J=2, NY
6270 DO 123 I=2, NX
6280 HEFF(I, J, 1)=HEFF(I, J, 1)+OUT1(I, J)*OUT2(I, J)
6290 123 CONTINUE
6300 CALL MEAN(AREA, OUT2)
6310 DO 124 J=2, NY
6320 DO 124 I=2, NX
6330 AREA(I, J, 1)=AREA(I, J, 1)+OUT1(I, J)*OUT2(I, J)
6340 124 CONTINUE
6350 CALL XSUM(HEFF1, THEFF1)
6351 THEFF1=THEFF1-THEFF
6360 * FIRST DO PREDICTOR
6370 DO 121 J=1, NY
6380 DO 121 I=1, NX
6390 UICE(I, J, 3)=UICE(I, J, 1)
6400 VICE(I, J, 3)=VICE(I, J, 1)
6410 UICEC(I, J)=UICE(I, J, 1)
6420 VICEC(I, J)=VICE(I, J, 1)
6430 121 CONTINUE
6440 * DO 121 J=1, NY
6450 * DO 121 I=1, NX
6460 * UICEC(I, J)=UICE(I, J, 1)
6470 * VICEC(I, J)=VICE(I, J, 1)
6480 *121 CONTINUE
6490 THETA=1. 0
6500 * DELTAT=DELT/2. 0
6510 DELTAT=DELT/2. 0
6520 CALL FORM(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, GAIIX, GAIY, GWATX, GWATY
&, FORCEX, FORCEY, HEFF, AMASS, AREA)
6530 CALL RELAX(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, AMASS, FORCEX, FORCEY
&, ERROR, THETA, UICEC, VICEC)
6560 * NOW DO REGULAR TIME STEP
6570 * NOW DOING BACKWARDS TIME STEP
6580 THETA=1. 0
6590 DELTAT=DELT
6600 CALL FORM(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, GAIIX, GAIY, GWATX, GWATY
&, FORCEX, FORCEY, HEFF, AMASS, AREA)
6620 * NOW SET U(1)=U(2) AND SAME FOR V
6630 DO 111 J=1, NY
6640 DO 111 I=1, NX
6650 UICE(I, J, 3)=UICE(I, J, 1)
6660 VICE(I, J, 3)=VICE(I, J, 1)
6670 UICEC(I, J)=UICE(I, J, 1)
6680 VICEC(I, J)=VICE(I, J, 1)
6690 * UICEC(I, J)=UICE(I, J, 1)
6700 * VICEC(I, J)=VICE(I, J, 1)
6710 UICE(I, J, 1)=UICE(I, J, 2)
6720 VICE(I, J, 1)=VICE(I, J, 2)
6730 111 CONTINUE
6740 CALL RELAX(UICE, VICE, ETA, ZETA, DRAGS, DRAGA, AMASS, FORCEX, FORCEY
&, ERROR, THETA, UICEC, VICEC)
6760 ICOUNT=ICOUNT+1
6770 * NOW DO ADVECTION
6780 CALL ADVECT(UICE, VICE, HEFF, DIFF1, LAD)
6790 CALL ADVECT(UICE, VICE, AREA, DIFF1, LAD)
6800 * NOW DO GROWTH
6810 CALL GROWTH(HEFF, AREA, HCDRR, HD, A22)
6820 * MUST CALL GROWTH ONLY AFTER CALLING ADVECTION
6830 * NOW CORRECT OUTFLOW PTS AND GET OUTFLOW ICE
6840 CALL XSUM(HEFF1, THEFF)
6850 DO 105 J=1, NY1
6860 DO 105 I=1, NX1
6870 HEFF(I, J, 1)=HEFF(I, J, 1)*OUT(I, J)

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6880      AREA(I, J, 1)=AREA(I, J, 1)*OUT(I, J)
6885      FHEFF(I, J)=FHEFF(I, J)*OUT(I, J)
6886      HDIFF(I, J)=(1.0-AREA(I, J, 2))*FO(I, J)*OUT(I, J)*DELT
6890 105    CONTINUE
6900      CALL XSUM(HEFF1, THEFF2)
6902      CALL XSUM(AREA1, ARSUM)
6903      CALL XSUM(FHE1, FHSUM)
6904      CALL XSUM(GR, GRSUM)
6905      GRSUM1=GRSUM1+GRSUM
6906      FHSUM1=FHSUM1+FHSUM
6907      ARSUM1=ARSUM1+ARSUM
6910      TOUT1=THEFF-THEFF2-THEFF1
6920      THEFF=THEFF2
6930      TOUT=TOUT+TOUT1
6935      ICOU=ICOUNT
6940 * NOW PRINT OUT EVERY KTH PT
6945 * IF(ICOUNT.LT.366) GO TO 97
6946 * ICOU=ICOUNT-365
6950      KTH=MOD(ICOUNT, 8)
6955      KTH=MOD(ICOUNT, 20)
6960      IF(KTH.EQ.1) GO TO 95
6970      GO TO 96
6980 95    CONTINUE
6990 4     FORMAT(//)
7000 1     FORMAT(1X, 'TIME STEP AND TOTAL THICKNESS ARE', 2G20.12)
7010      PRINT 1, ICOUNT, THEFF1
7020      PRINT 1, ICOUNT, THEFF
7030      PRINT 4
7040      PRINT 2, TOUT1, TOUT
7050      PRINT 4
7060 2     FORMAT(1X, 'OUTFLOW AND NET OUTFLOW ARE', 2G14.7)
7070      PRINT 3, ((UICE(I, J, 1), I=1, NX), J=1, NY)
7080      PRINT 4
7090      PRINT 3, ((VICE(I, J, 1), I=1, NX), J=1, NY)
7100      PRINT 4
7110      PRINT 3, ((HEFF(I, J, 1), I=1, NX), J=1, NY)
7120      PRINT 4
7130      PRINT 3, ((AREA(I, J, 1), I=1, NX), J=1, NY)
7140      PRINT 4
7150      PRINT 3, ((ZETA(I, J), I=1, NX), J=1, NY)
7160      PRINT 4
7170      PRINT 3, ((FHEFF(I, J), I=1, NX), J=1, NY)
7180      PRINT 4
7190      PRINT 3, ((STRESS(I, J, 1), I=1, NX), J=1, NY)
7200      PRINT 4
7210      PRINT 3, (FQROW(I), I=1, 21)
7220      PRINT 4
7230 3     FORMAT(1X, 5G20.14)
7240 * NOW PRINT OUT SQUARE VELOCITY AND VELOCITY DIFERENCE AND MAX CHANGE
7250      SQ1=0.0
7260      SQ=0.0
7270      DO 130 J=1, NY
7280      DO 130 I=1, NX
7290      SQ=SQ+UICE(I, J, 1)**2+VICE(I, J, 1)**2
7300      UERR(I, J)=UICE(I, J, 1)-UICE(I, J, 2)
7310      VERR(I, J)=VICE(I, J, 1)-VICE(I, J, 2)
7320      SQ1=SQ1+UERR(I, J)**2+VERR(I, J)**2
7330 130    CONTINUE
7340 * ISU=MAXMAG(UERRV)
7350 * ISV=MAXMAG(VERRV)
7355 * SMU=DABS(UERRV(ISU+1))
7356 * SMV=DABS(VERRV(ISV+1))
7360      CALL XMAXM(UERRV, SMU)
7370      CALL XMAXM(VERRV, SMV)
7380      SM=DMAX1(SMU, SMV)
7390      PRINT 5, SQ, SQ1, SM

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7400 5 FORMAT(1X, 'SQUARE VELOCITY, SG. VELOCITY DIFFERENCE, MAX CHANGE
7410 &/1X, 3G20. 12)
7420 PRINT 4
7430 96 CONTINUE
7440 * WRITE(20) ((UICE(I, J, 1), I=1, NX), J=1, NY)
7450 * WRITE(20) ((VICE(I, J, 1), I=1, NX), J=1, NY)
7460 * WRITE (20) ((HEFF(I, J, 1), I=1, NX1), J=1, NY1)
7470 * WRITE (20) ((AREA(I, J, 1), I=1, NX1), J=1, NY1)
7480 * WRITE (20) ((FHEFF(I, J, 1), I=1, NX1), J=1, NY1)
7490 * WRITE (20) ((ZETA(I, J), I=1, NX1), J=1, NY1)
7500 * WRITE (20) ((STRESS(I, J, 1), I=1, NX1), J=1, NY1)
7505 * WRITE(23) TOUT1, FHSUM, GRSUM, ARSUM, THEFF
7510 97 CONTINUE
7511 PRINT 1, ICOU, THEFF
7512 PRINT 2, TOUT1, TOUT
7513 PRINT 8, FHSUM, FHSUM1
7514 PRINT 9, GRSUM, GRSUM1, ARSUM, ARSUM1
7515 9 FORMAT(1X, 'OPEN WATER GROWTH AND NET GROWTH ETC', 4G15. 7)
7516 8 FORMAT(1X, 'GROWTH AND NET GROWTH ARE', 2G14. 7)
7520 * NOW CHECK COUNTER AND DECIDE IF NEW WINDS NEEDED
7530 KTH1=MOD(ICOUNT, 365)
7540 IF(KTH1.EQ.0) GO TO 211
7550 GO TO 212
7560 211 CONTINUE
7570 * REWIND 31
7580 PRINT 7
7590 7 FORMAT (1X, 'WIND FILE REWOUND')
7600 212 CONTINUE
7610 * READ(31) GAIRX, GAIRY
7620 * NOW DECIDE IF DONE
7630 IF(ICOUNT.GT.50) GO TO 200
7635 IF(ICOUNT.GT.20) GO TO 200
7640 GO TO 100
7650 200 CONTINUE
7660 KSTOP=2
7670 IF (KSTOP.EQ.1) GO TO 202
7680 GO TO 201
7690 202 CONTINUE
7700 DO 204 J=1, NY
7710 DO 204 I=1, NX
7720 GAIRX(I, J)=0. 0
7730 GAIRY(I, J)=0. 0
7740 204 CONTINUE
7750 ICOUNT=0
7760 KSTOP=2
7770 GO TO 100
7780 201 CONTINUE
7790 * WRITE(21) UICE, VICE, HEFF, AREA
7800 STOP
7810 END
7820 SUBROUTINE MEAN(HEFF, HMEAN)
7830 * SUBROUTINE FINDS MEAN HEFF AT OUTFLOW PTS OF VALUES AROUND
7840 IMPLICIT DOUBLE PRECISION (A-H, O-Z)
7850 PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
7860 &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
7870 DIMENSION HEFF(NX1, NY1, 3), HMEAN(NX1, NY1)
7880 COMMON/OUTFLO/OUT(NX1, NY1)
7890 DO 101 J=2, NY
7900 DO 101 I=2, NX
7910 HMEAN(I, J)=(HEFF(I+1, J, 1)*OUT(I+1, J)+HEFF(I+1, J+1, 1)*OUT(I+1, J+1)
7920 & +HEFF(I+1, J-1, 1)*OUT(I+1, J-1)+HEFF(I, J+1, 1)*OUT(I, J+1)
7930 & +HEFF(I, J-1, 1)*OUT(I, J-1)+HEFF(I-1, J, 1)*OUT(I-1, J)
7940 & +HEFF(I-1, J+1, 1)*OUT(I-1, J+1)+HEFF(I-1, J-1, 1)*OUT(I-1, J-1)
7950 & /(OUT(I+1, J)+OUT(I+1, J+1)+OUT(I+1, J-1)+OUT(I, J+1)+OUT(I, J-1)
7960 & +OUT(I-1, J)+OUT(I-1, J+1)+OUT(I-1, J-1)+. 00001D+00)
7970 101 CONTINUE

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7980      RETURN
7990      END
8000      SUBROUTINE PLAST(UICE, VICE, PRESS, ETA, ZETA, ECCEN, ZMAX, ZMIN)
8010 * SUBROUTINE CALCULATES STRAIN RATES AND VISCOUS PARAMETERS
8020      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
8030      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
8040      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
8050      DIMENSION UICE(NX, NY, 3), VICE(NX, NY, 3), PRESS(NX1, NY1)
8060      &, ETA(NX1, NY1), ZETA(NX1, NY1), E11(NX1, NY1), E22(NX1, NY1)
8070      &, E12(NX1, NY1), ZMAX(NX1, NY1), ZMIN(NX1, NY1)
8080      COMMON/STEP/DELTAT, DELTAX, DELTAY
8090      COMMON/STRESS/STRESS(NX1, NY1, 3)
8100      COMMON/ARRAY/HEFFM(NX1, NY1), UVM(NX, NY)
8110      ECM2=1. 0/(ECCEN**2)
8120      GMIN=1. 0D-20
8130 * NOW EVALUATE STRAIN RATES
8140      DO 101 J=2, NY
8150      DO 101 I=2, NX
8160      E11(I, J)=(0. 5/DELTAX)*(UICE(I, J, 1)+UICE(I, J-1, 1)
8170      & -UICE(I-1, J, 1)-UICE(I-1, J-1, 1))
8180      E22(I, J)=(0. 5/DELTAY)*(VICE(I, J, 1)+VICE(I-1, J, 1)
8190      & -VICE(I, J-1, 1)-VICE(I-1, J-1, 1))
8200      E12(I, J)=(0. 25/DELTAY)*(UICE(I, J, 1)+UICE(I-1, J, 1)
8210      & -UICE(I, J-1, 1)-UICE(I-1, J-1, 1))
8220      & +(0. 25/DELTAX)*(VICE(I, J, 1)+VICE(I, J-1, 1)
8230      & -VICE(I-1, J, 1)-VICE(I-1, J-1, 1))
8240 * NOW EVALUATE VISCOSITIES
8250      DELT=(E11(I, J)**2+E22(I, J)**2)*(1. 0+ECM2)+4. 0*ECM2*E12(I, J)**2
8260      & +2. 0*E11(I, J)*E22(I, J)*(1. 0-ECM2)
8270      DELT1=DSQRT(DELT)
8280      DELT1=DMAX1(GMIN, DELT1)
8290      ZETA(I, J)=0. 5*PRESS(I, J)/DELT1
8300 101  CONTINUE
8310 * NOW PUT MIN AND MAX VISCOSITIES IN
8320      DO 102 J=1, NY1
8330      DO 102 I=1, NX1
8340      ZETA(I, J)=DMIN1(ZMAX(I, J), ZETA(I, J))
8350      ZETA(I, J)=DMAX1(ZMIN(I, J), ZETA(I, J))
8360      ETA(I, J)=ECM2*ZETA(I, J)
8370      E11(I, J)=E11(I, J)*HEFFM(I, J)
8380      E22(I, J)=E22(I, J)*HEFFM(I, J)
8390      E12(I, J)=E12(I, J)*HEFFM(I, J)
8400      SS11=(ZETA(I, J)-ETA(I, J))*(E11(I, J)+E22(I, J))-PRESS(I, J)*0. 5
8410      STRESS(I, J, 1)=(2. 0*ETA(I, J)*E11(I, J)+SS11)
8420      STRESS(I, J, 2)=2. 0*ETA(I, J)*E22(I, J)+SS11
8430      STRESS(I, J, 3)=2. 0*ETA(I, J)*E12(I, J)
8440 102  CONTINUE
8450      RETURN
8460      END
9000      SUBROUTINE RELAX(OUICE, OVICE, OETA, OZETA, ODRAGS, ODRAGA
9010      &, OAMASS, ORCEX, ORCEY, ORROR, OTHETA, OUICEC, OVICEC)
9020      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
9030      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
9040      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
9050      DIMENSION UICE(NX, NY, 3), VICE(NX, NY, 3), ETA(NX1, NY1), ZETA(NX1, NY1)
9060      &, DRAGS(NX, NY), DRAGA(NX, NY), FORCEX(NX, NY), FORCEY(NX, NY), FXETA(4)
9070      &, FYZETA(4), FYETA(4), FYZETA(4), UERR(NX, NY), UERRV(N4)
9080      &, VERR(NX, NY), VERRV(N4), COEF(NX, NY), AMASS(NY, NY)
9090      &, FXM(NX, NY), FYM(NX, NY)
9100      &, UICEC(NX, NY), VICEC(NX, NY)
9110      &, FXE(4, NX, NY), FYE(4, NX, NY), FXZ(4, NX, NY), FYZ(4, NX, NY)
9120      &, COEFI(NX, NY)
9130      &, OUICE(NX, NY, 3), OVICE(NX, NY, 3), OETA(NX1, NY1), OZETA(NX1, NY1)
9140      &, ODRAGS(NX, NY), ODRAGA(NX, NY), OAMASS(NX, NY), ORCEX(NX, NY)
9150      &, ORCEY(NX, NY), OUICEC(NX, NY), OVICEC(NX, NY)
9160      &, HEFFM(NX1, NY1), UVM(NX, NY)

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9170      COMMON/DINV/DELIN2
9180      COMMON/STEP/OTAT, OTAX, OTAY
9190      COMMON/ARRAY/OHEFFM(NX1, NY1), OUVVM(NX, NY)
9200      EQUIVALENCE (UERR, UERRV), (VERR, VERRV)
9210      ICOUNT=0
9220      WFA=1. 5
9230 * NOW REPLACE DOUBLE PRECISION INPUT BY SINGLE PRECISION VARIABLES
9240      DO 601 J=1, NY
9250      DO 601 I=1, NX
9260      UVM(I, J)=OUVM(I, J)
9270      DRAGS(I, J)=ODRAGS(I, J)
9280      DRAGA(I, J)=ODRAGA(I, J)
9290      AMASS(I, J)=OAMASS(I, J)
9300      FORCEX(I, J)=ORCEX(I, J)
9310      FORCEY(I, J)=ORCEY(I, J)
9320      UICEC(I, J)=OUICEC(I, J)
9330      VICEC(I, J)=OVICEC(I, J)
9340      601 CONTINUE
9350      DO 603 K=1, 3
9360      DO 603 J=1, NY
9370      DO 603 I=1, NX
9380      UICE(I, J, K)=OUICE(I, J, K)
9390      VICE(I, J, K)=OVICE(I, J, K)
9400      603 CONTINUE
9410      DO 602 J=1, NY1
9420      DO 602 I=1, NX1
9430 *      HEFFM(I, J)=OHEFFM(I, J)
9440      ETA(I, J)=OETA(I, J)
9450      ZETA(I, J)=OZETA(I, J)
9460      602 CONTINUE
9470      ERROR=ORROR
9480      THETA=OTHETA
9490      DELTAX=OTAX
9500 *      DELTAY=OTAY
9510      DELTAT=OTAT
9520      DELIN =1. 0/DELTAX
9530      DELIN2=0. 5/(DELTAX**2)
9540      K=1
9550 * MUST UPDATE HEFF BEFORE CALLING RELAC
9560 * FIRST SET U(2)=U(1)
9570      DO 99 J=1, NY
9580      DO 99 I=1, NX
9590 * NOW MAKE SURE BDRY PTS ARE EQUAL TO ZERO
9600      UICE(I, J, 2)=UICE(I, J, 1)
9610      VICE(I, J, 2)=VICE(I, J, 1)
9620      UICE(I, J, 2)=OUICE(I, J, 1)
9630      OVICE(I, J, 2)=OVICE(I, J, 1)
9640      UICE(I, J, 1)=UICE(I, J, 3)*OUVM(I, J)
9650      OVICE(I, J, 1)=OVICE(I, J, 3)*OUVM(I, J)
9660      UICE(I, J, 1)=UICE(I, J, 3)*UVM(I, J)
9670      VICE(I, J, 1)=VICE(I, J, 3)*UVM(I, J)
9680      99 CONTINUE
9690 * NOW SET UP COEFFICIENTS OF DIAGONAL COMPONENTS
9700      501 CONTINUE
9710      DO 102 J=1, NY
9720      DO 102 I=1, NX
9730      COEF(I, J)=AMASS(I, J)/DELTAT+2. 0*THETA*(0. 5*DRAGS(I, J)
9740      &+2. 0*((ETA(I, J)+ETA(I+1, J)+ETA(I, J+1)+ETA(I+1, J+1))
9750      &+. 5*(ZETA(I, J)+ZETA(I+1, J)+ZETA(I, J+1)+ZETA(I+1, J+1))
9760      &)/(4. 0*(DELTAX**2)))
9770      COEF(I, J)=1. 0/COEF(I, J)
9780      102 CONTINUE
9790 * NOW CALCULATE ALL FUNCTIONS OF PREVIOUS U AND V VALUES
9800      TTTHETA=2. 0*(1. 0-THETA)
9810      DO 111 J=2, NYM1
9820      DO 111 I=2, NXM1

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9830 502 CONTINUE
9840 CALL FELLIP(UICE, VICE, ETA, FXETA, I, J, 2)
9850 CALL FELLIP(UICE, VICE, ZETA, FXZETA, I, J, 2)
9860 CALL FELLIP(VICE, UICE, ETA, FYETA, I, J, 2)
9870 CALL FELLIP(VICE, UICE, ZETA, FYZETA, I, J, 2)
9880 FXO=0. 5*(FXETA(1)+FXZETA(1)+FXETA(2)+FXETA(3)+FXZETA(4)-FXETA(4))
9890 FXO=TTTHETA*FXO
9900 FX1=(AMASS(I, J)/DELTAT-TTTHETA*0. 5*DRAGS(I, J))*UICE(I, J, 2)
9910 FX2=TTTHETA*0. 5*DRAGA(I, J)*VICE(I, J, 2)
9920 FYO=0. 5*(FYETA(1)+FYETA(2)+FYZETA(2)+FYZETA(3)-FYETA(3)+FYETA(4))
9930 FYO=FYO*TTTHETA
9940 FY1=(AMASS(I, J)/DELTAT-TTTHETA*0. 5*DRAGS(I, J))*VICE(I, J, 2)
9950 FY2=-TTTHETA*0. 5*DRAGA(I, J)*UICE(I, J, 2)
9960 FXC=AMASS(I, J)*0. 5*TTTHETA*
9970 & (UICEC(I, J)*(UICE(I+1, J, 2)-UICE(I-1, J, 2))
9980 & +VICEC(I, J)*(UICE(I, J+1, 2)-UICE(I, J-1, 2)))/(2. 0*DELTAX)
9990 FXM(I, J)=FXO+FX1+FX2+FORCEX(I, J)+FXC
10000 FYC=AMASS(I, J)*0. 5*TTTHETA*
10010 & (UICEC(I, J)*(VICE(I+1, J, 2)-VICE(I-1, J, 2))
10020 & +VICEC(I, J)*(VICE(I, J+1, 2)-VICE(I, J-1, 2)))/(2. 0*DELTAX)
10030 FYM(I, J)=FYO+FY1+FY2+FORCEY(I, J)+FYC
10040 111 CONTINUE
10050 * NOW SET U(3)=U(1)
10060 100 CONTINUE
10070 DO 101 J=1, NY
10080 DO 101 I=1, NX
10090 UICE(I, J, 3)=UICE(I, J, 1)
10100 VICE(I, J, 3)=VICE(I, J, 1)
10110 101 CONTINUE
10120 * NOW BEGIN SWEEP
10130 CALL FELLD1(UICE, VICE, ETA, FXE, 1)
10140 CALL FELLD1(UICE, VICE, ZETA, FXZ, 1)
10150 CALL FELLD1(VICE, UICE, ETA, FYE, 1)
10160 CALL FELLD1(VICE, UICE, ZETA, FYZ, 1)
10170 DO 103 J=2, NYM1
10180 DO 103 I=2, NXM1
10190 504 CONTINUE
10200 K=1
10210 FXETA(1) = FXE(1, I, J)+DELIN2*
10220 &(UICE(I-1, J, K)*( ETA(I, J+1)+ ETA(I, J)))
10230 FXETA(2)=FXE(2, I, J)+DELIN2*
10240 &(UICE(I, J-1, K)*(ETA(I, J)+ETA(I+1, J)))
10250 FXETA(3)=FXE(3, I, J)+0. 5*DELIN2*
10260 &(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*
10270 &(-ETA(I, J)+ETA(I+1, J))-VICE(I+1, J-1, K)*ETA(I+1, J)
10280 &+VICE(I-1, J, K)*(-ETA(I, J+1)+ETA(I, J))
10290 &-VICE(I-1, J+1, K)*ETA(I, J+1))
10300 FXETA(4)=FXE(4, I, J)+DELIN2*0. 5*
10310 &(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*
10320 &(-ETA(I+1, J)+ETA(I, J))-VICE(I+1, J-1, K)*ETA(I+1, J)
10330 &+VICE(I-1, J, K)*(ETA(I, J+1)-ETA(I, J))
10340 &-VICE(I-1, J+1, K)*ETA(I, J+1))
10350 FYETA(1)=FYE(1, I, J)+DELIN2*
10360 &(VICE(I-1, J, K)*( ETA(I, J+1)+ ETA(I, J)))
10370 FYETA(2)=FYE(2, I, J)+DELIN2*
10380 &(VICE(I, J-1, K)*(ETA(I, J)+ETA(I+1, J)))
10390 FYETA(3)=FYE(3, I, J)+0. 5*DELIN2*
10400 &(UICE(I-1, J-1, K)*ETA(I, J)+UICE(I, J-1, K)*
10410 &(-ETA(I, J)+ETA(I+1, J))-UICE(I+1, J-1, K)*ETA(I+1, J)
10420 &+UICE(I-1, J, K)*(-ETA(I, J+1)+ETA(I, J))
10430 &-UICE(I-1, J+1, K)*ETA(I, J+1))
10440 FYETA(4)=FYE(4, I, J)+DELIN2*0. 5*
10450 &(UICE(I-1, J-1, K)*ETA(I, J)+UICE(I, J-1, K)*
10460 &(-ETA(I+1, J)+ETA(I, J))-UICE(I+1, J-1, K)*ETA(I+1, J)
10470 &+UICE(I-1, J, K)*(ETA(I, J+1)-ETA(I, J))
10480 &-UICE(I-1, J+1, K)*ETA(I, J+1))

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10490      FXZETA(1)=FXZ(1, I, J)+DELIN2*
10500      &(UICE(I-1, J, K)*(ZETA(I, J+1)+ZETA(I, J)))
10510      FXZETA(4)=FXZ(4, I, J)+DELIN2*0. 5*
10520      &(VICE(I-1, J-1, K)*ZETA(I, J)+VICE(I, J-1, K)*
10530      &(-ZETA(I+1, J)+ZETA(I, J))-VICE(I+1, J-1, K)*ZETA(I+1, J)
10540      &+VICE(I-1, J, K)*(ZETA(I, J+1)-ZETA(I, J))
10550      &-VICE(I-1, J+1, K)*ZETA(I, J+1))
10560      FYZETA(2)=FYZ(2, I, J)+DELIN2*
10570      &(VICE(I, J-1, K)*(ZETA(I, J)+ZETA(I+1, J)))
10580      FYZETA(3)=FYZ(3, I, J)+0. 5*DELIN2*
10590      &(UICE(I-1, J-1, K)*ZETA(I, J)+UICE(I, J-1, K)*
10600      &(-ZETA(I, J)+ZETA(I+1, J))-UICE(I+1, J-1, K)*ZETA(I+1, J)
10610      &+UICE(I-1, J, K)*(-ZETA(I, J+1)+ZETA(I, J))
10620      &-UICE(I-1, J+1, K)*ZETA(I, J+1))
10630      FX3=0. 5*(FXETA(1)+FXZETA(1)+FXETA(2)+FXETA(3)+FXZETA(4)-FXETA(4))
10640      FX3=FX3*2. 0*THETA
10650      FXCP=AMASS(I, J)*THETA*
10660      & (UICEC(I, J)*(UICE(I+1, J, 1)-UICE(I-1, J, 1))
10670      &+VICEC(I, J)*(UICE(I, J+1, 1)-UICE(I, J-1, 1)))*0. 5*DELIN
10680      FX3=FX3+FXCP
10690      FY3=0. 5*(FYETA(1)+FYETA(2)+FYZETA(2)+FYZETA(3)-FYETA(3)+FYETA(4))
10700      FY3=FY3*2. 0*THETA
10710      FYCP=AMASS(I, J)*THETA*
10720      & (UICEC(I, J)*(VICE(I+1, J, 1)-VICE(I-1, J, 1))
10730      &+VICEC(I, J)*(VICE(I, J+1, 1)-VICE(I, J-1, 1)))*0. 5*DELIN
10740      FY3=FY3+FYCP
10750      FL11=0. 5*DRAGA(I, J)*COEFI(I, J)
10760      FL11=FL11*2. 0*THETA
10770      F11=(FXM(I, J)+FX3)*COEFI(I, J)
10780      F22=(FYM(I, J)+FY3)*COEFI(I, J)
10790 505 CONTINUE
10800      FL11S=1. 0+FL11**2
10810      FL11SI=1. 0/FL11S
10820      UICOR=((F11+FL11*F22)*FL11SI)*UVM(I, J)
10830      VICOR=((F22-FL11*F11)*FL11SI)*UVM(I, J)
10840      OVICE(I, J, 1)=OVICE(I, J, 1)+WFA*(UICOR-UICE(I, J, 1))
10850      OVICE(I, J, 1)=OVICE(I, J, 1)+WFA*(VICOR-VICE(I, J, 1))
10860 *     OVICE(I, J, 1)=OVICE(I, J, 1)+1. 5*(UICOR-UICE(I, J, 1))
10870 *     OVICE(I, J, 1)=OVICE(I, J, 1)+1. 5*(VICOR-VICE(I, J, 1))
10880      UICE(I, J, 1)=OVICE(I, J, 1)
10890      VICE(I, J, 1)=OVICE(I, J, 1)
10900 *     UICE(I, J, 1)=UICE(I, J, 1)+1. 5*(UICOR-UICE(I, J, 1))
10910 *     VICE(I, J, 1)=VICE(I, J, 1)+1. 5*(VICOR-VICE(I, J, 1))
10920 103 CONTINUE
10930      ICOUNT=ICOUNT+1
10940      IF(ICOUNT.GT.800) GO TO 201
10950      IF(ICOUNT.LT.100) GO TO 202
10960      WFA=1. 0
10970 202 CONTINUE
10980 * NOW CHECK MAX ERROR
10990 *     S11=0. 0
11000 *     S22=0. 0
11010 * FORM ERROR MATRIX
11020      DO 104 J=1,NY
11030      DO 104 I=1,NX
11040      UERR(I, J)=UICE(I, J, 1)-UICE(I, J, 3)
11050      VERR(I, J)=VICE(I, J, 1)-VICE(I, J, 3)
11060 104 CONTINUE
11070 * NOW FIND ERROR
11080      CALL XMAXM(UERRV, S1)
11090      CALL XMAXM(VERRV, S2)
11100 *     S1=ABS(UERRV(IU+1))
11110 *     S2=ABS(VERRV(IV+1))
11120      S1=DMAX1(S1, S2)
11130      IF(S1.LT.ERROR) GO TO 200

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11140      GO TO 100
11150  201  CONTINUE
11160      PRINT 11
11170  11  FORMAT(IX, 'NO CONVERGENCE AFTER 800 ITERATIONS')
11180 * NOW END
11190  200  CONTINUE
11200 *      PRINT 1, ICOUNT
11210      PRINT 12, S1
11220      PRINT 1, ICOUNT
11230  12  FORMAT(IX, 'MAX ERROR AND U AND V POWER', 3G12. 5)
11240  1  FORMAT(IX, 'NO OF ITERATIONS ARE', 1G12. 5)
11250      RETURN
11260      END
11270      SUBROUTINE FELLIP(UICE, VICE, ETA, F, I, J, K)
11280 * SPACER
11290      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
11300      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
11310      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
11320      DIMENSION UICE(NX, NY, 3), VICE(NX, NY, 3), ETA(NX1, NY1), F(4)
11330      COMMON/STEP/DTAT, DTAX, DTAY
11340      DELTAX=DTAX
11350      S1=. 5/(DELTAX**2)
11360      F(1)=S1*(UICE(I+1, J, K)*(ETA(I+1, J+1)+ETA(I+1, J))
11370      &-UICE(I, J, K)*(ETA(I+1, J+1)+ETA(I, J)+ETA(I+1, J)+ETA(I, J+1))
11380      &+UICE(I-1, J, K)*(ETA(I, J+1)+ETA(I, J)))
11390      F(2)=S1*(UICE(I, J+1, K)*(ETA(I+1, J+1)+ETA(I, J+1))
11400      &-UICE(I, J, K)*(ETA(I+1, J+1)+ETA(I, J)+ETA(I+1, J+1))
11410      &+UICE(I, J-1, K)*(ETA(I, J)+ETA(I+1, J)))
11420      F(3)=S1*(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*(-ETA(I, J)
11430      &+ETA(I+1, J))-VICE(I+1, J-1, K)*ETA(I+1, J)+VICE(I-1, J, K)*(-ETA(I, J+1)
11440      &+ETA(I, J))+VICE(I, J, K)*(-ETA(I, J)-ETA(I+1, J+1)+ETA(I+1, J)
11450      &+ETA(I, J+1)))
11460      F(3)=F(3)+S1*(VICE(I+1, J, K)*(-ETA(I+1, J)+ETA(I+1, J+1))
11470      &-VICE(I-1, J+1, K)*ETA(I, J+1)
11480      &+VICE(I, J+1, K)*(-ETA(I+1, J+1)+ETA(I, J+1))
11490      &+VICE(I+1, J+1, K)*ETA(I+1, J+1))
11500      F(4)=S1*(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*(-ETA(I+1, J)
11510      &+ETA(I, J))-VICE(I+1, J-1, K)*ETA(I+1, J)+VICE(I-1, J, K)*(-ETA(I, J+1)
11520      &-ETA(I, J))+VICE(I, J, K)*(ETA(I, J+1)+ETA(I+1, J)-ETA(I, J)-ETA(I+1,
11530      &J+1)))
11540      F(4)=F(4)+S1*(VICE(I+1, J, K)*(ETA(I+1, J)-ETA(I+1, J+1))
11550      &-VICE(I-1, J+1, K)*ETA(I, J+1)
11560      &+VICE(I, J+1, K)*(ETA(I+1, J+1)-ETA(I, J+1))
11570      &+VICE(I+1, J+1, K)*ETA(I+1, J+1))
11580      F(3)=F(3)*. 5
11590      F(4)=F(4)*. 5
11600      RETURN
11610      END
11620      SUBROUTINE FELLD1(UICE, VICE, ETA, F, K)
11630 * SPACER
11640      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
11650      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
11660      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
11670      DIMENSION UICE(NX, NY, 3), VICE(NX, NY, 3), ETA(NX1, NY1), F(4, NX, NY)
11680 *      COMMON/STEP/DELTAT, DELTAX, DELTAY
11690      COMMON/DINV/S1
11700      DO 101 J=2, NYM1
11710      DO 101 I=2, NXM1
11720      F(1, I, J)=S1*(UICE(I+1, J, K)*(ETA(I+1, J+1)+ETA(I+1, J)))
11730      F(2, I, J)=S1*(UICE(I, J+1, K)*(ETA(I+1, J+1)+ETA(I, J+1)))
11740      F(3, I, J)=S1*(VICE(I, J, K)*(-ETA(I, J)-ETA(I+1, J+1)+ETA(I+1, J)
11750      &+ETA(I, J+1))
11760      &+VICE(I+1, J, K)*(-ETA(I+1, J)+ETA(I+1, J+1))
11770      &+VICE(I, J+1, K)*(-ETA(I+1, J+1)+ETA(I, J+1))
11780      &+VICE(I+1, J+1, K)*ETA(I+1, J+1))
11790      F(4, I, J)=S1*(VICE(I, J, K)*(ETA(I, J+1)+ETA(I+1, J)-ETA(I, J)-ETA(I+1,

```

```

11800      &J+1))+VICE(I+1, J, K)*(ETA(I+1, J)-ETA(I, J+1))
11810      &+VICE(I, J+1, K)*(ETA(I+1, J+1)-ETA(I, J+1))
11820      &+VICE(I+1, J+1, K)*ETA(I+1, J+1))
11830      F(3, I, J)=0. 5*F(3, I, J)
11840      F(4, I, J)=0. 5*F(4, I, J)
11850 101  CONTINUE
11860      RETURN
11870      END
11880      SUBROUTINE FELLD(VICE, VICE, ETA, F, I, J, K, FETA)
11890 * SPACER
11900      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
11910      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
11920      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
11930      DIMENSION UICE(NX, NY, 3), VICE(NX, NY, 3), ETA(NX1, NY1), F(4)
11940      &, FETA(4, NX, NY)
11950      COMMON/DINV/S1
11960      F(1)=FETA(1, I, J)+S1*(
11970      &+UICE(I-1, J, K)*(ETA(I, J+1)+ETA(I, J)))
11980      F(2)=FETA(2, I, J)+S1*((
11990      &+UICE(I, J-1, K)*(ETA(I, J)+ETA(I+1, J)))
12000      F(3)=FETA(3, I, J)+0. 5*S1*(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*
12010      &(-ETA(I, J))
12020      &+ETA(I+1, J))-VICE(I+1, J-1, K)*ETA(I+1, J)+VICE(I-1, J, K)*(-ETA(I, J+1)
12030      &+ETA(I, J))
12040      &-VICE(I-1, J+1, K)*ETA(I, J+1))
12050      F(4)=FETA(4, I, J)+S1*0. 5*
12060      &(VICE(I-1, J-1, K)*ETA(I, J)+VICE(I, J-1, K)*(-ETA(I+1, J)
12070      &+ETA(I, J))-VICE(I+1, J-1, K)*ETA(I+1, J)+VICE(I-1, J, K)*(ETA(I, J+1)
12080      &-ETA(I, J))
12090      &-VICE(I-1, J+1, K)*ETA(I, J+1))
12100      RETURN
12110      END
12120      SUBROUTINE XMAXM(UERRV, X)
12130 *
12140      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
12150      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
12160      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1, N4=NX*NY
12170      DIMENSION UERRV(N4)
12180      X=0. 0
12190      DO 100 I=1, N4
12200      X=DMAX1(DABS(UERRV(I)), X)
12210 100  CONTINUE  ..
12220      RETURN
12230      END
12240      SUBROUTINE XSUM(HEFF1, S1)
12250 * PROGRAM SUMS UP VECTOR
12260      IMPLICIT DOUBLE PRECISION (A-H, O-Z)
12270      PARAMETER NX=5, NY=6, N3=(NX+1)*(NY+1), NX1=NX+1
12280      &, NY1=NY+1, NXM1=NX-1, NYM1=NY-1
12290      DIMENSION HEFF1(N3)
12300      S1=0. 0
12310      DO 100 I=1, N3
12320      S1=S1+HEFF1(I)
12330 100  CONTINUE
12340      RETURN
12350      END

```

APPENDIX B: TWENTY-ONE-DAY TEST RUN OF MODEL

This 21-day test run is obtained by compiling and executing the code as listed in Appendix A. The only input data needed are contained in a file BNDATA which is read by subroutine BNDRY (lines 1020-1050). This BNDATA file consists of the following formatted data which define the boundary conditions shown in Figure 2:

```
00000001110011100111001110000000
000000011110011110011110011110000000
00000001110001111001111001111001111000000
```

INTEGER VARIABLE (K2) SET, NEVER USED, LINES 250 AND 300

DOUBLE PRECISION VARIABLE (DWAT) SET, NEVER USED, LINE 2660
 DOUBLE PRECISION VARIABLE (DAIR) SET, NEVER USED, LINE 2670.
 DOUBLE PRECISION VARIABLE (GRAV) SET, NEVER USED, LINE 2750.
 DOUBLE PRECISION VARIABLE (EPSIL) SET, NEVER USED, LINE 3210.
 DOUBLE PRECISION VARIABLE (ETAC) SET, NEVER USED, LINE 3240.

DOUBLE PRECISION VARIABLE (ONE) SET, NEVER USED, LINE 3890.
 DOUBLE PRECISION VARIABLE (ZERO) SET, NEVER USED, LINE 3900.

INTEGER VARIABLE (KGRO) SET, NEVER USED, LINE 6200

00000000000000	00000000000000	.00000000000000	00000000000000	00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	00000000000000	00000000000000	00000000000000	00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	00000000000000	00000000000000	00000000000000	00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	00000000000000	00000000000000	00000000000000	00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	1 00000000000000	1 00000000000000	1 00000000000000	1 00000000000000
00000000000000	00000000000000	00000000000000	00000000000000	00000000000000
HESUM IS 19 0000	UVSUM IS 12 00000			
HESUM IS 20.0000	UVSUM IS 12 00000			
3 20000000000000	3 30000000000000	3. 40000000000000	3 50000000000000	3 60000000000000
3 30000000000000	3 40000000000000	3 50000000000000	3 60000000000000	3 70000000000000
3 40000000000000	3 50000000000000	3 60000000000000	3 70000000000000	3 80000000000000
3 50000000000000	3 60000000000000	3 70000000000000	3 80000000000000	3 90000000000000
3 60000000000000	3 70000000000000	3 80000000000000	3 90000000000000	4 00000000000000
3 70000000000000	3 80000000000000	3 90000000000000	4 00000000000000	4 10000000000000
-3 60000000000000	-3 40000000000000	-3 20000000000000	-3 00000000000000	-2 80000000000000
-3 40000000000000	-3 20000000000000	-3 00000000000000	-2 80000000000000	-2 60000000000000
-3 20000000000000	-3 00000000000000	-2 80000000000000	-2 60000000000000	-2 40000000000000
-3 00000000000000	-2 80000000000000	-2 60000000000000	-2 40000000000000	-2 20000000000000
-2 80000000000000	-2 60000000000000	-2 40000000000000	-2 20000000000000	-2 00000000000000
-2 60000000000000	-2 40000000000000	-2 20000000000000	-2 00000000000000	-1 80000000000000
10000000000000E-01	90000000000000E-02	80000000000000E-02	70000000000000E-02	60000000000000E-02
11000000000000E-01	10000000000000E-01	90000000000000E-02	80000000000000E-02	70000000000000E-02
12000000000000E-01	11000000000000E-01	10000000000000E-01	90000000000000E-02	80000000000000E-02
13000000000000E-01	12000000000000E-01	11000000000000E-01	10000000000000E-01	90000000000000E-02
14000000000000E-01	13000000000000E-01	12000000000000E-01	11000000000000E-01	10000000000000E-01
15000000000000E-01	14000000000000E-01	13000000000000E-01	12000000000000E-01	11000000000000E-01

1 . 139930555556E-05
 2 . 225694444444E-06
 3 . 532407407407E-07
 4 . 428240740741E-07
 5 . 358796296296E-07
 6 . 312500000000E-07
 7 . 243055555556E-07
 8 . 162037037037E-07
 9 . 104166666667E-07
 10 . 347222222222E-08
 11 . 000000000000
 12 . -231481481481E-08
 13 . -347222222222E-08
 14 . -462962962963E-08
 15 . -578703703704E-08
 16 . -578703703704E-08
 17 . -694444444444E-08
 18 . -694444444444E-08
 19 . -694444444444E-08
 20 . -694444444444E-08
 21 . -694444444444E-08

- 91984855243125E-02 . 16909133238402E-01 . 15793329942552E-01 . 53719427407723E-01 . 25985353244302E-01
 - 48514339437489E-01 . 44866731198317E-02 . 32137460399924E-02 . 41095469141595E-01 . 52528457534129E-01
 - 49100120038751E-01 . 37432349310818E-02 . 25349005658436E-02 . 13850900150383E-02 . 78931630880927E-01
 - 49613531169529E-01 . 30748993469716E-02 . 19335467202533E-02 . 85320645905588E-03 . 78690221589046E-01
 - 50051513532798E-01 . 24843979837337E-02 . 14120133532538E-02 . 40264520065144E-03 . 78531100243467E-01
 - 11202822487627E-01 . 14676039471344E-01 . 13893139793508E-01 . 13174723471373E-01 . 52167493728903E-01

- 31270146475486E-01 . 61104461249715E-01 . 59131530183871E-01 . 25013964552678E-01 . 92762427753814E-02
 56249348874907E-02 . 97804294854293E-02 . 11471236315805E-01 . 20727131162579E-01 . 20541965411089E-01
 84022932127701E-02 . 12713525336701E-01 . 14350815038782E-01 . 15952311038522E-01 . 15776677375776E-01
 11133006149244E-01 . 15609595065236E-01 . 17203577711480E-01 . 18772625921528E-01 . 18356824528410E-01
 13626766612019E-01 . 18478933637849E-01 . 20040342450109E-01 . 21588086253953E-01 . 20943549621534E-01
 47818393763294E-01 . 84200415473009E-01 . 85959404535894E-01 . 87716287349946E-01 . 55311079804659E-01

MAX ERROR AND U AND V POWER . 31432E-05

NO OF ITERATIONS ARE 11.000

. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000
. 00000000000000	. 55123024654173E-03	. 82694462070983E-03	. 12280766662109E-02	. 00000000000000
. 00000000000000	. 39393391967157E-03	. 4999519458721E-03	. 29665577472785E-03	. 00000000000000
. 00000000000000	. 19239739927943E-03	. 27581610960687E-03	. 19194841853472E-03	. 00000000000000
. 00000000000000	. -14394697336957E-05	. 12611643083897E-03	. 19914970340528E-03	. 00000000000000
. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000

. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000
. 00000000000000	. 64468343209846E-03	. 70332404696872E-03	. 93073828893523E-04	. 00000000000000
. 00000000000000	. 95417502989914E-03	. 11875137686298E-02	. B4244398417197E-03	. 00000000000000
. 00000000000000	. 10528211003079E-02	. 14008863763407E-02	. 11776913529051E-02	. 00000000000000
. 00000000000000	. 79048475252198E-03	. 10379124877481E-02	. 93301620599902E-03	. 00000000000000
. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000	. 00000000000000

MAX ERROR AND U AND V POWER . 23423E-05

NO OF ITERATIONS ARE 11.000

MAX ERROR AND U AND V POWER . 40834E-05

NO OF ITERATIONS ARE 8.0000

TIME STEP AND TOTAL THICKNESS ARE 1.000000000000

3.29669230773

TIME STEP AND TOTAL THICKNESS ARE 1.000000000000

62.6692788514

OUTFLOW AND NET OUTFLOW ARE .9147826E-04 .9147826E-04

.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.41057247550924E-04	.56846035029838E-04	.81468058298827E-04	.000000000000000
.000000000000000	.28138928200678E-04	.33823664003268E-04	.18482411011226E-04	.000000000000000
.000000000000000	.12451366106910E-04	.18646878541670E-04	.13083451735438E-04	.000000000000000
.000000000000000	-.17442721549091E-05	.93035439167671E-05	.14553120213194E-04	.000000000000000
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.54421502682481E-04	.61766223432941E-04	.13906653062935E-05	.000000000000000
.000000000000000	.82072462056398E-04	.10502405617843E-03	.74681432747438E-04	.000000000000000
.000000000000000	.90320025744148E-04	.12079027872319E-03	.10016494703124E-03	.000000000000000
.000000000000000	.66972170170303E-04	.89820544754459E-04	.79701598118132E-04	.000000000000000
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	3.2982791312153	3.2982375791387	3.2982877200253	.000000000000000
.000000000000000	3.2982775716448	3.2982826576149	3.2982446106205	3.2984181030719
.000000000000000	3.2983322686140	3.2983470141199	3.2983647286222	3.2983948719973
.000000000000000	3.2984023133110	3.2984301493117	3.2984468665240	3.2984427139725
.000000000000000	3.2984662040978	3.2985539676780	3.2985750780206	3.2984953018081
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.99996700313530	.99718959907456	1.000000000000000	.000000000000000
.000000000000000	.9999653006558	.99996807280986	.99442693180934	1.000000000000000
.000000000000000	.99998312147957	.99998759428303	.99999296768207	.99723731117466
.000000000000000	1.000000000000000	1.000000000000000	1.000000000000000	1.000000000000000
.000000000000000	1.000000000000000	1.000000000000000	1.000000000000000	1.000000000000000
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	4120879120879.1	4120879120879.1	4120879120879.1	.000000000000000
.000000000000000	4120879120879.1	4120879120879.1	3910418581531.3	4120879120879.1
.000000000000000	4120879120879.1	4120879120879.1	4120879120879.1	4120879120879.1
.000000000000000	4120879120879.1	4120879120879.1	4120879120879.1	4120879120879.1
.000000000000000	4120879120879.1	4120879120879.1	3630562648686.7	4120879120879.1
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02	.000000000000000
.000000000000000	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02
.000000000000000	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02
.000000000000000	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02
.000000000000000	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02	.16846153846154E-02
.000000000000000	.000000000000000	.000000000000000	.000000000000000	.000000000000000
.000000000000000	-3941.0269520577	-2998.2341796075	-4615.0152605708	-6808461394898.3
.000000000000000	-3338.6697984968	-4740.3650366613	-2960.4275986860	-11643.622254355
.000000000000000	-5745.5721754968	-7692.1710279809	-8373.7826942921	-9389.7083890993
.000000000000000	-7943.2686226044	-8559.3741244123	-9082.1036741092	-10038.926290415
.000000000000000	-10397.702305097	-13449.670173992	-13633.964519457	-11813.305386092
.1399305555556E-03	.223674444444444E-06	.53240740740741E-07	.42824074074074E-07	.39879629629630E-07
.31250000000000E-07	.2430555555556E-07	.16203703703704E-07	.10416666666667E-07	.34722222222222E-08
.000000000000000	-.23148148148148E-08-	.34722222222222E-08-	.46296296296296E-08-	.37870370370370E-08
-.57870370370370E-08-	.694444444444444E-08-	.694444444444444E-08-	.694444444444444E-08-	.694444444444444E-08
-.69444444444444E-08				

SQUARE VELOCITY, SQ. VELOCITY DIFFERENCE, MAX CHANGE
 .966152836119E-07 .119455001931E-04 .128009609762E-02

TIME STEP AND TOTAL THICKNESS ARE 1.00000000000 62. 6692788514
 OUTFLOW AND NET OUTFLOW ARE .9147826E-04 .9147826E-04
 GROWTH AND NET GROWTH ARE .3200769E-01 .3200769E-01
 OPEN WATER GROWTH AND NET GROWTH ETC .0000000 .0000000 18. 98872 18. 98872
 MAX ERROR AND U AND V POWER .41442E-05
 NO OF ITERATIONS ARE 2.0000
 MAX ERROR AND U AND V POWER .24269E-05
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 2.00000000000 62. 7024471339
 OUTFLOW AND NET OUTFLOW ARE .8751464E-04 .1789929E-03
 GROWTH AND NET GROWTH ARE .3325580E-01 .6526349E-01
 OPEN WATER GROWTH AND NET GROWTH ETC .1363857E-02 .1363857E-02 18. 99109 37. 97980
 MAX ERROR AND U AND V POWER .28397E-05
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .13899E-05
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 3.00000000000 62. 7353028718
 OUTFLOW AND NET OUTFLOW ARE .8244857E-04 .2614415E-03
 GROWTH AND NET GROWTH ARE .3293819E-01 .982016BE-01
 OPEN WATER GROWTH AND NET GROWTH ETC .1077774E-02 .2441631E-02 18. 99294 56. 97274
 MAX ERROR AND U AND V POWER .96390E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .88624E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 4.00000000000 62. 7679033279
 OUTFLOW AND NET OUTFLOW ARE .7928826E-04 .3407297E-03
 GROWTH AND NET GROWTH ARE .3267974E-01 .1308814
 OPEN WATER GROWTH AND NET GROWTH ETC .8536515E-03 .3295283E-02 18. 99438 75. 96712
 MAX ERROR AND U AND V POWER .66153E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .67121E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 5.00000000000 62. 8002938478
 OUTFLOW AND NET OUTFLOW ARE .7855221E-04 .4192819E-03
 GROWTH AND NET GROWTH ARE .3246907E-01 .1633505
 OPEN WATER GROWTH AND NET GROWTH ETC .6795474E-03 .3974830E-02 18. 99549 94. 96261
 MAX ERROR AND U AND V POWER .46274E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .45260E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 6.00000000000 62. 8325112472
 OUTFLOW AND NET OUTFLOW ARE .7921190E-04 .4984938E-03
 GROWTH AND NET GROWTH ARE .3229661E-01 .1956471
 OPEN WATER GROWTH AND NET GROWTH ETC .5454454E-03 .4520276E-02 18. 99634 113. 9590
 MAX ERROR AND U AND V POWER .36399E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .31219E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 7.00000000000 62. 8645845271
 OUTFLOW AND NET OUTFLOW ARE .8060446E-04 .57909B3E-03
 GROWTH AND NET GROWTH ARE .3215388E-01 .2278010
 OPEN WATER GROWTH AND NET GROWTH ETC .4424561E-03 .4962732E-02 18. 99699 132. 9559
 MAX ERROR AND U AND V POWER .28442E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .21506E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 8.00000000000 62. 8965363328
 OUTFLOW AND NET OUTFLOW ARE .8238654E-04 .6614848E-03
 GROWTH AND NET GROWTH ARE .3203419E-01 .2598352
 OPEN WATER GROWTH AND NET GROWTH ETC .3635552E-03 .5326287E-02 18. 99748 151. 9534
 MAX ERROR AND U AND V POWER .22296E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .14870E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 9.00000000000 62. 9283854123
 OUTFLOW AND NET OUTFLOW ARE .8433947E-04 .7458243E-03
 GROWTH AND NET GROWTH ARE .3193342E-01 .2917686
 OPEN WATER GROWTH AND NET GROWTH ETC .3044336E-03 .5630721E-02 18. 99785 170. 9513
 MAX ERROR AND U AND V POWER .17618E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .10368E-06
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 10.00000000000 62. 9601454288
 OUTFLOW AND NET OUTFLOW ARE .8631673E-04 .8321410E-03
 GROWTH AND NET GROWTH ARE .3184633E-01 .3236149
 OPEN WATER GROWTH AND NET GROWTH ETC .2596003E-03 .5890321E-02 18. 99813 189. 9494
 MAX ERROR AND U AND V POWER .14327E-06

NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .74802E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 11.0000000000 62.9918267462
 OUTFLOW AND NET OUTFLOW ARE .8824076E-04 .9203818E-03
 GROWTH AND NET GROWTH ARE .3176956E-01 .3553845
 OPEN WATER GROWTH AND NET GROWTH ETC .2255131E-03 .6115834E-02 18.99835 208.9478
 MAX ERROR AND U AND V POWER .11812E-06
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .60845E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 12.0000000000 63.0234373316
 OUTFLOW AND NET OUTFLOW ARE .9006798E-04 .1010450E-02
 GROWTH AND NET GROWTH ARE .3170065E-01 .3870851
 OPEN WATER GROWTH AND NET GROWTH ETC .1996125E-03 .6315447E-02 18.99851 227.9463
 MAX ERROR AND U AND V POWER .99636E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .53976E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 13.0000000000 63.0549833090
 OUTFLOW AND NET OUTFLOW ARE .9177519E-04 .1102225E-02
 GROWTH AND NET GROWTH ARE .3163775E-01 .4187229
 OPEN WATER GROWTH AND NET GROWTH ETC .1799415E-03 .6495388E-02 18.99864 246.9449
 MAX ERROR AND U AND V POWER .85949E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .49112E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 14.0000000000 63.0864693874
 OUTFLOW AND NET OUTFLOW ARE .9335328E-04 .1195578E-02
 GROWTH AND NET GROWTH ARE .3157943E-01 .4503023
 OPEN WATER GROWTH AND NET GROWTH ETC .1650067E-03 .6660395E-02 18.99873 265.9436
 MAX ERROR AND U AND V POWER .80681E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .45665E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 15.0000000000 63.1178991909
 OUTFLOW AND NET OUTFLOW ARE .9480144E-04 .1290380E-02
 GROWTH AND NET GROWTH ARE .3152460E-01 .4818269
 OPEN WATER GROWTH AND NET GROWTH ETC .1536697E-03 .6814064E-02 18.99880 284.9424
 MAX ERROR AND U AND V POWER .76632E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .43203E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 16.0000000000 63.1492755131
 OUTFLOW AND NET OUTFLOW ARE .9612368E-04 .1386503E-02
 GROWTH AND NET GROWTH ARE .3147245E-01 .5132994
 OPEN WATER GROWTH AND NET GROWTH ETC .1450639E-03 .6959128E-02 18.99885 303.9413
 MAX ERROR AND U AND V POWER .73288E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .41433E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 17.0000000000 63.1806005122
 OUTFLOW AND NET OUTFLOW ARE .9732675E-04 .1483830E-02
 GROWTH AND NET GROWTH ARE .3142233E-01 .5447217
 OPEN WATER GROWTH AND NET GROWTH ETC .1385303E-03 .7097658E-02 18.99889 322.9402
 MAX ERROR AND U AND V POWER .69765E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .40019E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 18.0000000000 63.2118759303
 OUTFLOW AND NET OUTFLOW ARE .9842014E-04 .1582250E-02
 GROWTH AND NET GROWTH ARE .3137384E-01 .5760955
 OPEN WATER GROWTH AND NET GROWTH ETC .1336429E-03 .7231301E-02 18.99892 341.9391
 MAX ERROR AND U AND V POWER .67617E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .38332E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 19.0000000000 63.2431039050
 OUTFLOW AND NET OUTFLOW ARE .9942083E-04 .1681671E-02
 GROWTH AND NET GROWTH ARE .3132700E-01 .6074225
 OPEN WATER GROWTH AND NET GROWTH ETC .1304179E-03 .7361719E-02 18.99894 360.9381
 MAX ERROR AND U AND V POWER .67242E-07
 NO OF ITERATIONS ARE 1.0000
 MAX ERROR AND U AND V POWER .38222E-07
 NO OF ITERATIONS ARE 1.0000
 TIME STEP AND TOTAL THICKNESS ARE 20.0000000000 63.2742841446
 OUTFLOW AND NET OUTFLOW ARE .1003302E-03 .1782001E-02

GROWTH AND NET GROWTH ARE . 3128097E-01 6387035
 OPEN WATER GROWTH AND NET GROWTH ETC 1279838E-03 7489703E-02 18 99896 379. 9370
 MAX ERROR AND U AND V POWER . 66773E-07
 NO OF ITERATIONS ARE 1. 0000
 MAX ERROR AND U AND V POWER 36318E-07
 NO OF ITERATIONS ARE 1. 0000
 TIME STEP AND TOTAL THICKNESS ARE 21. 0000000000 3 32958600935
 TIME STEP AND TOTAL THICKNESS ARE 21. 0000000000 63 3054187372

OUTFLOW AND NET OUTFLOW ARE . 0111530E-03 1883154E-02

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. 000000000000000	. -9211. 6345199173	. -10009. 352354911	. -10193. 336151179	. -9556. 6844731167

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. 00000000000000E-08 . 23148148148148E-08- . 34722222222222E-08- . 46296296296294E-08- . 57870370370370E-08
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. 69444444444444E-08

SQUARE VELOCITY, SQ. VELOCITY DIFFERENCE, MAX CHANGE
. 893987030642E-07 . 658437122999E-13 . 847259876693E-07

TIME STEP AND TOTAL THICKNESS ARE . 21. 0000000000 . 63. 3054187372
OUTFLOW AND NET OUTFLOW ARE . 1011530E-03 . 1883154E-02
GROWTH AND NET GROWTH ARE . 3123575E-01 . 6699393
OPEN WATER GROWTH AND NET GROWTH ETC . 1263262E-03 . 7616029E-02 . 18. 99897 . 398. 9360

14. 277 sec. 501 I/O

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